

Resolving the chlorine isotope composition of Earth’s depleted upper mantle by examining the effects of mantle heterogeneity

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The chlorine isotope ratio ($\delta^{37}\text{Cl}$) of Earth’s depleted mantle has implications for volatile exchange between Earth’s surface and mantle. However, estimates vary greatly, ranging from $< -3\text{‰}$ to $+0.9\text{‰}$ [1, 2, 3, 4]. Previous studies did not fully consider the possible impact of mantle heterogeneity on the $\delta^{37}\text{Cl}$ values of measured mid-ocean ridge basalt (MORB) glasses. To evaluate possible mantle heterogeneity, we measured the $\delta^{37}\text{Cl}$ values via IRMS and SIMS on 30 MORB glasses from several ridge segments spanning a wide range of Cl/K (0.01-0.82), trace element (Figure 1A), and radiogenic isotope (Figure 1B) compositions. By examining variations in $\delta^{37}\text{Cl}$ values with indices of subduction influence and shallow assimilation, we are able to better determine the depleted MORB-source mantle (DMM) $\delta^{37}\text{Cl}$ value and the impacts of subduction on mantle $\delta^{37}\text{Cl}$ values.

Measured MORB $\delta^{37}\text{Cl}$ values range from -1.5‰ to $+1.0\text{‰}$. Samples affected by shallow assimilation (high Cl/K) extend towards $\delta^{37}\text{Cl}$ values lower than seawater, suggesting assimilation of hydrothermal brines does not necessarily result in seawater-like $\delta^{37}\text{Cl}$ values. After filtering for shallow assimilation, the most depleted MORB samples from several localities have $\delta^{37}\text{Cl}$ values of -0.4 to -0.5‰ (Figure 1). For all localities, samples with lower $^{143}\text{Nd}/^{144}\text{Nd}$ and higher $\text{K}_2\text{O}/\text{TiO}_2$ extend to higher and/or lower $\delta^{37}\text{Cl}$ values than the most depleted MORB samples (Figure 1). This likely reflects incorporation of subducted material (altered oceanic crust and marine sediments) with heterogenous $\delta^{37}\text{Cl}$ values into the DMM with a $\delta^{37}\text{Cl}$ value of -0.4‰ to -0.5‰ (Figure 1). Much of the disagreement between previous estimates of the DMM $\delta^{37}\text{Cl}$ value can be attributed to the effects of shallow assimilation and mantle heterogeneity.

[1] Sharp et al. (2007), Nature 446, 1062-1065. [2] Bonifacie et al. (2008), Science 319, 1518-1520. [3] Layne et al. (2009), Geol. 37(5), 427-430. [4] Pinti et al. (2020), Geochim. Cosmochim Acta 276, 14-30. [5] John et al. (2010), Earth Planet. Sci. Lett. 298, 175-182. [6] Halldórsson et al. (2016), Geol. 44(8), 679-682.

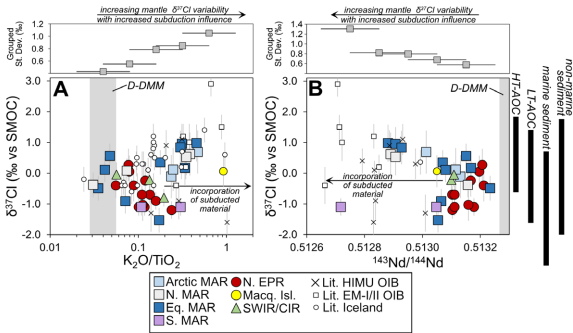


Figure 1. Chlorine isotope ratios of MORB samples from this study and MORB and OIB samples from literature filtered for Cl contamination plotted against (A) $\text{K}_2\text{O}/\text{TiO}_2$ and (B) $^{143}\text{Nd}/^{144}\text{Nd}$. Error bars are 2 SD. Colored symbols include MORB data from this study and from previous studies [1, 2]. Literature OIB data are from [5, 6]. Gray vertical bars are the most depleted endmember of DMM $\text{K}_2\text{O}/\text{TiO}_2$ and $^{143}\text{Nd}/^{144}\text{Nd}$. The black bars to the right show the range of $\delta^{37}\text{Cl}$ values in subducted materials (HT-AOC= high-temperature altered oceanic crust, LT-AOC= low-temperature altered oceanic crust). Grouped St. Dev.= grouped standard deviations, which shows the standard deviation in $\delta^{37}\text{Cl}$ values of a subset of MORB and OIB samples for a range of $\text{K}_2\text{O}/\text{TiO}_2$ (left) or $^{143}\text{Nd}/^{144}\text{Nd}$ (right) shown by the horizontal black lines. N. MAR= Northern Mid-Atlantic Ridge (MAR), Eq. MAR= Equatorial MAR, S. MAR= Southern MAR, N. EPR= Northern East Pacific Rise, Macq. Isl.= Macquarie Island, SWIR= Southwest Indian Ridge, CIR= Central Indian Ridge, Lit= Literature, OIB= Ocean Island Basalt.