

# The study of non-hydrocarbon gases and its implications for natural hydrogen and helium exploration

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Natural gases may contain significant amounts of non-hydrocarbon components such as N<sub>2</sub>, CO<sub>2</sub>, and H<sub>2</sub>S. The prediction of non-hydrocarbon contents, which may range from traces to nearly 100%, remains one of the major challenges in energy gas exploration. Many aspects regarding origins and accumulation mechanism of nonhydrocarbon gases in deep subsurface reservoirs are still debated. The present state of understanding is summarized in Table 1.

The discovery of surface emanations of natural gases rich in H<sub>2</sub>, has prompted speculations about commercial reservoirs of “natural hydrogen”, and several mechanisms have been proposed.

## 1. Hydrogen (H<sub>2</sub>) generation by serpentinization of olivine:

This is a viable and active process that has persistently fuelled chemoautotrophic life throughout much of Earth's history. The reaction of water with ferrous iron consumes oxygen and releases H<sub>2</sub>. Most known occurrences are in Mid-Ocean hydrothermal areas and surface-near water-ophiolite environments, which typically lack efficient seals and structural traps.

## 2. Thermal liberation of H<sub>2</sub> from sedimentary organic matter (disproportioning):

Sedimentary organic matter carries its “own” hydrogen, deriving from water splitting and “H-fixation” during photosynthesis. Laboratory experiments show that large amounts of H<sub>2</sub> are released during *open* system pyrolysis. But only marginal quantities are observed in *closed* system pyrolysis, due to the high reactivity of hydrogen.

## 3. Liberation of H<sub>2</sub> during thermal decomposition of ammonium clays/feldspars:

Ammonium clays/feldspars are the most likely sources of N<sub>2</sub>-rich gases in the deep subsurface. The ammonium hydrogen is derived from water splitting during nitrogen fixation by *nitrogenases*. Thermal/hydrothermal decomposition of NH<sub>4</sub><sup>+</sup>-bearing minerals producing N<sub>2</sub> should release significant quantities of H<sub>2</sub>. Bleached hematite layers in the vicinity of N<sub>2</sub>-rich gas reservoirs may result from reduction of ferric iron by H<sub>2</sub> or NH<sub>4</sub><sup>+</sup>. Data from N<sub>2</sub>/CO<sub>2</sub>-rich natural gases in German Lower Permian reservoirs document H<sub>2</sub> contents of up to 4%. These

were, however, attributed to corrosion and have never been examined systematically.

## Correlation of He and N<sub>2</sub> contents of natural gases:

Helium contents of natural gases appear to correlate more clearly with N<sub>2</sub> contents than with concentrations of other gases. This could indicate a common source but also similar dependencies on temperature and burial history.

Table 1: Proposed sources of non-hydrocarbon natural gases

Gas	Proposed sources/origin:
N <sub>2</sub>	“Inorganic” (organic origin): ammonium-rich clay minerals
	Organic: pyrrolic/pyridinic organic compounds
	Inorganic: nitrates (hypothetical)
	Inorganic: mantle (hypothetical)
CO <sub>2</sub>	Inorganic: carbonate metamorphism and hydrolysis
	Organic: thermal decarboxylation, thermochemical sulfate reduction
H <sub>2</sub> S	Organic: “Bacterial” Sulfate Reduction (BSR)
	Inorganic: Thermochanical Sulfate Reduction (TSR)
H <sub>2</sub>	Inorganic: Serpentinization (water splitting by oxidation of ferrous iron)
	Organic: Decomposition of organic matter (disproportioning)
	Organic/inorganic: Decomposition of ammonium clay minerals