

Shocked minerals in the Northwest Africa 807 meteorite: Investigating the formation conditions of the high-pressure fayalite polymorph asimowite

MYRTO SIMOPOULOU¹, IOANNIS BAZIOTIS¹, JINPING HU², CHI MA², LUDOVIC FERRIÈRE^{3,4}, DIMITRIOS PALLES⁵ AND PAUL ASIMOW²

¹Agricultural University of Athens

²California Institute of Technology

³Natural History Museum Abu Dhabi, Abu Dhabi, United Arab Emirates

⁴Natural History Museum Vienna

⁵National Hellenic Research Foundation

Pressure and temperature conditions for shock synthesis of high-pressure (HP) polymorphs of olivine in meteorites are presumed to resemble those where these phases appear in the mantles of terrestrial planets [1], although the timescale and mechanisms of shock and static compression are different [1, 2]. Because the Fe-analogue of wadsleyite, asimowite [3], appears in shocked meteorites but is absent from equilibrium phase diagrams, its existence poses a challenge to calibration of shock pressure using static compression and to understanding of dynamic synthesis of HP olivine polymorphs. Our observations of HP phases in shock-induced melt veins (MVs) in NWA 807 [4] offer insight into the origin of Fe-rich wadsleyite and asimowite.

Eight HP phases -ringwoodite, ahrensite, wadsleyite, asimowite, majorite, albitic jadeite, lingunite and xieite- were verified by chemical analysis, EBSD and/or Raman methods. One lens-shaped MV clast (**Fig. 1**) hosts majorite ($\text{Ca}_{0.08}\text{Mg}_{2.62}\text{Fe}_{0.99}\text{Mn}_{0.08}\text{Al}_{0.40}\text{Si}_{3.75}\text{O}_{12}$) at the center, rimmed by wadsleyite ($\text{Fa}_{30.3}$), followed by ringwoodite (Fa_{38-41}), an extended domain of olivine ($\text{Fa}_{12.2}$), and an outer rim of ahrensite (Fa_{51-54}). At two locations in the MV (one shown in **Fig. 2**), we identify the third known occurrence of asimowite (Fa_{50-53}), coexisting with ahrensite (Fa_{76}).

The HP phases in NWA 807 suggest pressures from 14 to 23 GPa and temperature heterogeneities during the shock [2]. Based on $\text{Mg}_2\text{SiO}_4\text{-Fe}_2\text{SiO}_4$ phase relations at 1600 °C, we infer that asimowite (Fa_{53}) and ahrensite (Fa_{76}) represent a metastable phase coexistence within the fayalite stability field at ~8 GPa [5]. This suggests that asimowite crystallizes metastably from ahrensite during decompression, with transformation into fayalite kinetically hindered or barricaded by high activation-energy. This potentially explains why asimowite is found in meteorites despite its absence from equilibrium phase diagrams and suggests a need for rapid decompression experiments or simulations to study the kinetics of phase transformation in the olivine system.

[1] Gruetzner *et al.* (2018) *Earth Planet Sci Lett* **482**, 236-244.
 [2] Hu & Sharp (2022) *Prog Earth Planet Sci* **9**, 1-22. [3] Bindi

et al. (2019) *Am Mineral* **104**, 775-778. [4] Simopoulou *et al.* (2024) *LPI Contrib.* 3036, 6054. [5] Fei *et al.* (1991) *J Geophys Res: Solid Earth* **96**, 2157-2169.

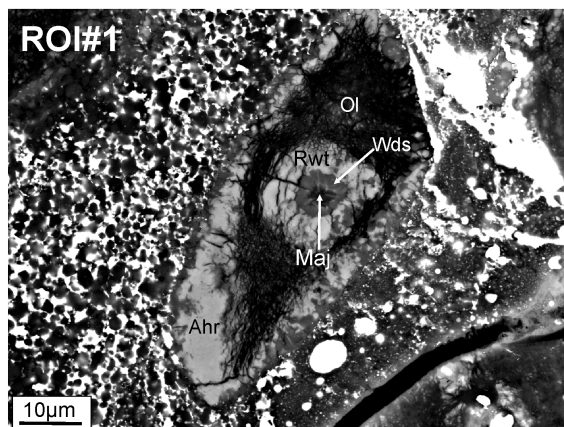


Figure 1: BSE image of the lens-shaped clast showing various HP zones (ROI#1). Ahr: ahrensite, rwt: ringwoodite, wds: wadsleyite, ol: olivine, maj: majorite garnet

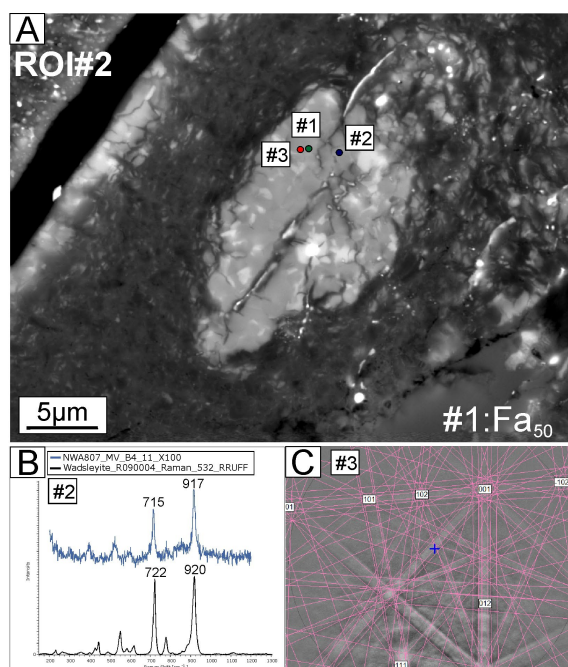


Figure 2: Wadsleyite clast containing asimowite confirmed by chemical analysis, Raman, and EBSD (ROI#2)