

Abiotic mineral-fluid-gas reactivity in underground H₂ storage: A kinetic rate law for reductive pyrite dissolution from 60-150 °C and up to 150 bar P_{H2}

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Hydrogen is an energy storage medium as excess energy from renewable sources can be used to produce green H₂. Large-scale H₂ storage in subsurface environments (e.g. salt caverns, depleted gas fields, aquifers) can support the energy transition. However, fluid-rock-H₂ interactions are a concern as H₂ is a potent electron donor. For example, the reductive dissolution of pyrite produces gaseous H₂S, a toxic and corrosive contaminant of the stored H₂ gas [1, 2].

To study the kinetics of reductive pyrite dissolution at conditions relevant to underground H₂ storage, anoxic alteration experiments of pyrite ($\phi = 50\text{-}100\ \mu\text{m}$) were conducted at $T = 60\text{-}150\ ^\circ\text{C}$, $P_{\text{H}_2} = 7\text{-}150\ \text{bar}$ and $pH \sim 3.8\text{-}8.6$ in 30 mmol/L NaCl solutions. Experiments were carried out in H₂-pressurized batch reactors made either of titanium or coated stainless steel. The temporal evolution of dissolved sulfide concentrations was measured by methylene blue spectrophotometry (Fig. 1). Electron microscopy of post-mortem pyrite grains revealed the formation of authigenic pyrrhotite and magnetite. The derived kinetic rate law was based on total (aqueous + gas) sulfide production rates determined at 12h, which represents far-from-equilibrium conditions. For implementation in geochemical codes, the rate law was based on the general form given in Eq. (1).

Reductive pyrite dissolution was upscaled over a 30-year period using batch models to simulate H₂ storage into a sandstone depleted gas field, containing 0.13 vol.% pyrite, 4.6 vol.% carbonates at 117 °C and an injection pressure of 150 bar (PHREEQC, modified Thermoddem database [3]). The models predicted buffered gas concentrations of 28 ppm H₂S and 4.4 % CO₂. In a sensitivity study, calculated gas concentrations varied on the order of 10⁰-10³ ppm H₂S, emphasizing the role of Fe-bearing minerals (e.g., chlorite, siderite) as a source of Fe²⁺_(aq) for scavenging dissolved sulfide ions.

References:

[1] Bourgeois et al. (1979), *Rev. Inst. Franç. Pétrole* 34, 371-386.

[2] Truche et al. (2010), *Geochim. Cosmochim. Acta* 74, 2894-2914.

[3] van Winden et al. (2024), 5th EAGE-GET conference abstract.

$$\text{Equation (1): } R = Ae^{-E_a/RT}(a_{\text{H}^+})^{-n}(P_{\text{H}_2})^m\left(1 - \frac{Q}{K}\right)$$

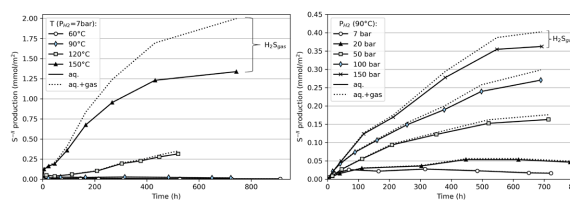


Fig. 1: Normalized sulfide production (mmol/m²) as a function of time, temperature (left) and P_{H2} (right). Aqueous sulfide concentrations were measured by methylene blue spectrophotometry while gaseous H₂S was computed in equilibrium.