

Magnesium- and iron-isotope evidence for the origin of kimberlitic melts from Udachnaya-East pipe (Siberian Craton, Russia)

WENYU WANG^{1,2}, ALEXANDER GOLOVIN³, DR. KATE KISEEVA^{2,4}, XIN DING¹, ZHENWU CHEN⁵ AND FANG HUANG¹

¹University of Science and Technology of China

²University College Cork

³Russian Academy of Sciences

⁴American Museum of Natural History

⁵Anhui University of Science and Technology

Kimberlite is a carbonate-rich volcanic rock representing deep mantle melts (>150 km). However, its post-emplacement alteration (e.g., serpentinisation) and a large amount of external material (xenoliths and xenocrysts) complicate the identification of the origin of primary kimberlite melt as well as the parental source to kimberlite magmas [1,2]. To address the origin of kimberlite magmas, we analysed major and trace elements, as well as Mg-Fe isotopes, from five rare, fresh, non-serpentinised kimberlites from Udachnaya-East, Siberian Craton, Russia. In addition to that, we analyse serpentinised kimberlites from the same pipe and sedimentary bed rock.

The Mg isotopic compositions ($\delta^{26}\text{Mg} = -0.31 \sim -0.24\text{‰}$) of non-serpentinized kimberlites fall within the range of the mantle ($\sim -0.25 \text{‰}$), MORB ($\sim -0.25 \text{‰}$) and OIB ($\sim -0.26 \text{‰}$), whereas their Fe isotopic compositions ($\delta^{56}\text{Fe} = 0.06 \sim 0.12 \text{‰}$) are heavier than the mantle ($\sim 0.00 \text{‰}$) and comparable to MORB ($\sim 0.10 \text{‰}$) and OIB ($\sim 0.11 \text{‰}$). Crustal contamination, fractional crystallisation, and diffusion processes are excluded as controls on these isotopic variations. By subtracting the contribution of mantle-derived components, we obtain the Mg-Fe isotopic compositions of the kimberlite melt ($\delta^{26}\text{Mg} = -0.34 \sim -0.24 \text{‰}$, $\delta^{56}\text{Fe} = 0.10 \sim 0.21 \text{‰}$). Kimberlite melts exhibit carbonate-like signatures, such as low Hf/Hf* (0.27~0.29) and Ti/Ti* (0.56~0.89), however, their Fe isotopic composition contrasts with that of igneous carbonates ($-0.30 \sim -0.10\text{‰}$) and aligns with mantle-derived silicate melts (0.05 ~ 0.33‰). The Mg-Fe isotopic range of kimberlite melts overlaps with that reported for melilitites ($\delta^{26}\text{Mg} = -0.27 \pm 0.04 \text{‰}$, $\delta^{56}\text{Fe} = 0.13 \pm 0.06\text{‰}$) [3,4], which are assumed to be generated by low-degree (1–2%) melting of carbonated peridotite.

We conclude that the isotopic data supports the origin of kimberlite magmas from low-degree melts of carbonated peridotite metasomatised by deep mantle carbonates, with source oxygen fugacity ranging from FMQ-2.9 to FMQ-0.3.

[1] Russell, J.K., et al., 2012, *Nature*, V. 481, P. 352-356. [2] Kjarsgaard, B., et al., 2009, *Lithos*, V. 112, P. 236-248. [3] Li, W.Y., et al. 2016, *EPSL*, V. 444, P. 26-33. [4] Weyer, S., et al., 2012, *ChemGeo*, V. 294, P. 42-50.