

Deglacial CO₂ release from the Southern Ocean during Termination IV

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Abrupt atmospheric CO₂ increases characterise a critical feature of deglaciations. The deglacial CO₂ rise toward Marine Isotope Stage (MIS) 9e (Termination IV) started from 197.1 ppm to 300.7 ppm at 335 ka BP^[1], representing the highest natural atmospheric CO₂ recorded in the Antarctic ice cores over the past 800 ka^[2]. Oceanic carbon storage changes must be involved in regulating the Pleistocene atmospheric CO₂ variations. However, the mechanisms and pathways of the air-sea carbon exchanges remain elusive partly due to the lack of oceanic carbonate system proxy data with a robust age control beyond Termination I. Here, we present high-resolution carbonate system records for Termination IV from Iberian Margin. We employ a new air-sea CO₂ exchange tracer ($[\text{CO}_3^{2-}]_{\text{as}}$)^[3] to reconstruct carbon transfer between the oceanic and atmospheric reservoirs. An increased $[\text{CO}_3^{2-}]_{\text{as}}$ would reflect enhanced CO₂ outgassing. At the onset of HS10.1, $[\text{CO}_3^{2-}]$ decreased rapidly, likely due to an expansion of southern-sourced Glacial Antarctic Bottom Water (GAABW). During the mid-HS10.1, the lack of any $[\text{CO}_3^{2-}]$ decline probably implies an increase in $[\text{CO}_3^{2-}]$ associated with the GAABW due to enhanced ventilation via the Southern Ocean. The increased $[\text{CO}_3^{2-}]_{\text{as}}$ indicated a net release of CO₂ from the Atlantic sector of the Southern Ocean during the mid-HS10.1. The centennial-scale atmospheric CO₂ rise at the end of the HS10.1 is unlikely related to the Atlantic sector of the Southern Ocean, as a $[\text{CO}_3^{2-}]_{\text{as}}$ decline was observed. Our results suggested that during the deglacial CO₂ increase toward MIS9e, there was a net release of CO₂ from the Atlantic sector of the Southern Ocean. The net CO₂ release is highly likely due to the expansion and ventilation of the GAABW.

References:

[1] Nehrbass-Ahles, C. et al. (2020), *Science* vol. 369 1000–1005.

[2] Bereiter, B. et al. (2015), *Geophys. Res. Lett.* **42**, 542–549.

[3] Yu, J. et al. (2019), *Nat. Commun.* **10**.