

## Setting upper limits on the natural hydrogen generation in intracratonic areas

KANCHANA KULARATNE<sup>1,2</sup>, PASCALE SÉNÉCHAL<sup>3</sup>,  
VALENTINE COMBAUDON<sup>2,4</sup>, OLIVIER SISSMANN<sup>4</sup>,  
ERIC DEVILLE<sup>4</sup> AND HANNELORE DERLUYN<sup>1,2</sup>

<sup>1</sup>Université de Pau et des Pays de l'Adour, E2S UPPA, CNRS, DMEX, Pau

<sup>2</sup>Université de Pau et des Pays de l'Adour, E2S UPPA, CNRS, TotalEnergies, LFCR, Pau

<sup>3</sup>Université de Pau et des Pays de l'Adour, E2S UPPA, CNRS, DMEX

<sup>4</sup>IFP Energies Nouvelles

Presenting Author: [kanchana.kularatne@univ-pau.fr](mailto:kanchana.kularatne@univ-pau.fr)

Most proposed exploration guidelines for natural H<sub>2</sub> exploration in intracratonic areas focus on geodynamics<sup>[1]</sup>, seismics<sup>[1]</sup>, mineralogical analysis<sup>[1]</sup> and mapping large scale H<sub>2</sub> seepages<sup>[2]</sup>. Here we present a complementary method of evaluating H<sub>2</sub> production potential based on X-Ray micro-computed tomography (micro-CT) and multi-modal imaging of the host-rocks. To obtain a first order estimation of H<sub>2</sub> generation, this method assumes that H<sub>2</sub> is generated via complete oxidation of Fe(II) in the Fe-rich minerals within the host-rock. This method neither considers the actual geochemical reaction/s that could yield different Fe(II):H<sub>2</sub> ratios nor the reaction kinetics. Micro-CT imaging combined with  $\mu$ Raman analysis is used to identify the Fe(II)-rich minerals in the imaged 3D micro-CT volume. However, one can also use quantitative electron microprobe (EMPA) maps or energy dispersive scanning electron microscopy (SEM EDS) maps to replace  $\mu$ Raman analysis, depending on the size of the sample to be analyzed. Once the Fe(II)-rich minerals are identified 2D micro-CT slice, the chemical and mineralogical information can be propagated within the imaged 3D micro-CT volume to obtain the volumes of Fe(II)-rich phases. Then, using the mineral density and theoretical or actual Fe(II) content in each phase, potential H<sub>2</sub> production in the host-rock can be obtained in H<sub>2</sub> (g)/host-rock (kg). Alternatively, this number can be calculated by obtaining the volumes using X-Ray powder diffraction (XRD) of the sample, but, the advantage here is the ability of analyzing a large sample such as a whole drill-core of the host-rock, non-destructively. Latest developments in spectral CT (Sp-CT) imaging allows better segmentation of Fe(II)-rich silicates, thus, allowing to model actual H<sub>2</sub>-producing reactions as well. This method could be used to set an upper limit on H<sub>2</sub> generation in intracratonic ophiolites, gabbros and banded iron formations reported in [3], [4] and [5].

[1] Dugamin et al (2019) ISRN Geonum-NST, 1, 16. [2] Moretti et al (2021) Geosciences, 11(3), 145. [3] Milcov (2022) Earth-Science Reviews, 230, 104063. [4] Combaudon et al (2023) Goldschmidt abstract #18169. [5] Geymond et al (2022) Minerals 2022, 12(2).