Determination of the oxidation state of arc primary melts using two fO₂ proxies

MARION GABORIEAU¹, **MURIEL LAUBIER**¹, NATHALIE BOLFAN-CASANOVA¹ AND MASSIMO POMPILIO²

¹Laboratoire Magmas et Volcans

²INGV-Sezione di Pisa

Presenting Author: muriel.laubier@uca.fr

Although many studies have demonstrated that arc magmas are more oxidized than mid-ocean ridge (MORB) and oceanic island basalts (OIB) [1-5], the oxidation state of their mantle source is still debated. This ongoing debate is mainly due to contradictory fO_2 values obtained from different proxies (e.g., Fe³⁺/\SigmaFe of olivine-hosted melt inclusions and glasses, Zn/ Σ Fe, V/Sc, V/Ga of lavas).

Here, we estimate the oxygen fugacity of high-Mg olivinehosted melt inclusions from mid-ocean ridges and various arcs (Aoba, Mount Meager, Vulcano, Stromboli), from one hot spot (Reunion Island) and Mount Etna using two fO₂ proxies: the $Fe^{3+}/\Sigma Fe$ of melts and the partition coefficient of V between olivine and melt (D_v^{Ol/Melt} [6]). Our results indicate that secondary processes such as H₂O and fO₂ re-equilibration, volatile degassing and fractional crystallization are unlikely to be responsible for the large variation in magmatic fO_2 . After reconstructing primary melt compositions, we show that (1) fO_2 values derived from $Fe^{3+}/\Sigma Fe$ and $D_v^{Ol/Melt}$ are comparable and (2) arc and Mount Etna primary melts are more oxidized than primary melts from mid-ocean ridges and Reunion Island. We then demonstrate, using trace element ratios, that the observed variability in primary melt fO2 is not due to the chemical variability of the mantle source prior to its metasomatism (using Zr/Nb) but is rather a consequence of the addition of slab-derived material (using Ba/Th and Sr/Th) during subduction processes.

[1] Kelley & Cottrell (2009), *Science* **325**, 605–607. [2] Kelley & Cottrell (2012), *Earth Planet. Sci. Lett.* **329–330**, 109–121. [3] Brounce et al. (2014), J. Pet. **55**(12), 2513–2536. [4] Brounce, et al. (2015), *Geology* **43**(9), 775–778. [5] Gaborieau et al. (2020), *Chem. Geol.* **547**, 119646. [6] Mallmann & O'Neill (2013), *J. Petrol.* **54**, 933-949.