

## The solubility of $\text{ErPO}_4$ in hydrothermal fluids at varying pH and salinity between 350 and 450°C

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Heavy rare earth elements (HREE) are becoming increasingly important due to their use in modern technologies [1]. HREE mobility in hydrothermal fluids is controlled by mineral stability [2] and aqueous speciation at elevated temperatures [3]. Thermodynamic data for some aqueous REE species are extrapolated from low temperatures [4], producing large discrepancies between experimentally derived and predicted xenotime solubilities [2]. This greatly limits our predictive capabilities to model REE transport in high temperature magmatic-hydrothermal fluids. Here we used batch-type Inconel 625 Parr reactors to investigate synthetic  $\text{ErPO}_4$  solubility between 350 and 450°C at 500 bar, varying pH (2, 3, 4, and 10), and salinities (0.01, 0.1, and 0.5 molal NaCl). At 350°C, xenotime solubility in acidic and alkaline pH solutions is higher with a depression proximal to pH 4.5 (Fig. 1). Aqueous Er concentrations are  $47.1 \pm 5.9$  ppt for pH 2,  $17.1 \pm 2.3$  ppt for pH 3,  $4.3 \pm 2.9$  ppt for pH 4, and  $16.1 \pm 4.2$  ppt for pH 10. Measured Er concentrations were compared to simulated Er speciation and total predicted Er solubility using the GEMS code package [5] and existing thermodynamic data [2,4]. Measured Er concentrations are 2-3.5 orders of magnitude lower versus simulated solubility using aqueous REE chloride and hydroxyl species from the Suprt92 dataset [4] (Fig. 1). Er speciation is controlled by chloride complexes (i.e.,  $\text{ErCl}^{2+}$ ) at acidic pH and by hydroxyl complexes in alkaline pH (i.e.,  $\text{Er(OH)}_3^0$ ,  $\text{Er(OH)}_4^-$ ). Simulated and experimental results contrast due to extrapolation of low temperature thermodynamic data for xenotime or aqueous Er chloride and hydroxyl species to 350°C. Previous work at lower temperatures resolved such discrepancies by adjusting the thermodynamic properties of xenotime [2]. Similar adjustments will need to be made with the data from this study. These obstacles highlight the unreliability of high temperature extrapolations and the need for high temperature experimental data.

[1] Migdisov et al., (2016), *Chem. Geol.* **439**, 13-42.

[2] Gysi et al., (2015), *Chem. Geol.* **401**, 83–95

[3] Migdisov and Williams-Jones, (2006) *Chem. Geol.* **234**, 17–27

