Quantitative Modeling of Magma Ocean in the early Earth and its implication for the origin of LLSVPs

DR. SEEMA KUMARI, PH.D.¹ AND RAMANANDA CHAKRABARTI²

¹Indian Institute of Science, Bangalore

²Indian Institute of Science

Presenting Author: seema091089@gmail.com

We modeled the evolution of terrestrial Magma Ocean (*MO*) using the short-lived ¹⁸²Hf-¹⁸²W isotope system ($t_{1/2}$ = 8.9 Myr) in a multi-reservoir Earth evolutionary model (Fig. 1). Model simulates several scenarios of *MO* crystallization. After crystallization, denser magma ocean residue (*MO_{RES}*) became gravitationally unstable and sank to the core-mantle boundary (CMB). The sequestrated *MO_{RES}* might represent the present-day seismically identified Large Low-Shear Velocity Provinces (LLSVPs) [1], and material from such a source is argued to have contributed to ocean island basalts (OIBs) [2].

We start with a chondritic Earth ($M_E = 5.9 \times 10^{24}$ kg) at $T_0 =$ 4.56 Ga. Core segregates and attains its present-day mass (1.883×10²⁴ kg), W concentration (500±120 ppb), ¹⁸²W/¹⁸⁴W ratio ($\mu^{182}W_{Core}$ of -220; $\mu^{182}W$ is ppm deviation in $^{182}W/^{184}W$ relative to present-day terrestrial standard) at the end of core formation (T_{CE}) . The time and rate of core formation constrains the starting composition of the silicate Earth (adopted from [3]) and impose several initialization states for MO models. We assume a large MO with a composition similar to the silicate Earth was formed soon after core formation, and crystallization of magma ocean, MO_{CX} occurs in time T_{CX} , leaving behind the residue, MO_{RES} ; $MO_{CX} + MO_{RES} = M_{MO}$. A series of differential equations describing the changing mass and abundance of ¹⁸²Hf, ¹⁸²W and ¹⁸⁴W nuclides in each reservoir are solved numerically at 0.1 Myr time steps for the initial 200 Myr of Earth's history. A successful solution reproduces the present-day Hf and W concentrations and ¹⁸²W/¹⁸²W ratios in MO_{RES} similar to those constrained from global OIBs, assuming OIBs originate from LLSVPs. Our modeling results constrain: (i) the amount of crystallization in the MO, MO_{CX} (ii) the timing of MOcrystallization, T_{CX} , and (iii) the depth of MO, d_{MO} . The composition of the model-derived MO_{RES} provides significant clues for early-formed distinct geochemical domains in the deep mantle that remained preserved until today and sampled by the deep-rising plumes.

[1] Ballmer, Lourenço, Hirose, Caracas & Nomura (2017), Geochem. Geophys. Geosyst. 18, 2785–2806.

[2] Garnero, McNamara & Shim (2016), Nat. Geosci. 9, 481-489.

[3] Kumari, Stracke, & Paul (2022), Chem. Geol. 611, 121104.

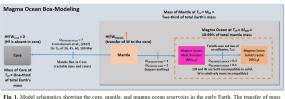


Fig. 1. Model schematics showing the core, mantle, and magma ocean reservoirs in the early Earth. The transfer of mass and species fluxes are shown by arrows, for example, an arrow from the human te to the core indicates core segregation from the bulk Earth.