

Evidence of colder ambient mantle condition during the Archean; Implication from the calculation of mantle potential temperature (Tp)

MR. SARBAJIT DASH, SENIOR RESEARCH SCHOLAR^{1,2}, EVSSK BABU^{1,2} AND JÉRÔME GANNE³

¹CSIR-National Geophysical Research Institute, Hyderabad, Uppal Road, 500007, India

²Academy of Scientific and Innovative Research (AcSIR), Ghaziabad-201002, India

³GET-OMP, Université de Toulouse, UPS, CNRS, IRD, CNES, 14 avenue E. Belin, 31400 Toulouse, France

Presenting Author: sarbajitdash1994@gmail.com

Mafic magmas indicate the physicochemical conditions e.g., pressure, temperature, and fluid availability during melt generation and its evolution in the continental or oceanic lithosphere. Here we present the mantle potential temperature (Tp) values from Indian cratons, assuming mantle melting and generation of mafic melt by keeping the redox condition $Fe^{+2}/Fe_{Total}=0.9$ at the magma source [1] and compared the result with previously published Tp values across the globe [2]. The calculated Tp values (1628°C-1525°C) during the Archean eventually overlap with that of the previously published data. However, the ca. 2.8 Ga basalt [3] from Mauranipur-Babina greenstone belt of the Bundelkhand craton yield a significantly lower Tp value (1473°C) than previously published literature. Trace element ratios e.g., La/Yb, Th/Ta, Nb/Y, and Zr/Y indicate MORB or Oceanic Plateau Basalt (OPB) signature, suggesting a depleted mantle source of origin for the samples. Similarly, in a Ta-Tb-Th triangular plot, the samples show the tholeiitic basalt signature. Aulbach and Arndt (2019) [4] calculated a Tp value <1500°C for the protoliths of mantle eclogite. The lack of direct evidence of low Tp producing basaltic crusts during the Archean from other cratons might indicate sample alteration, sampling biases and limitations in Tp calculation methodology. However, the direct evidence of lower Tp from this study indicates a relatively colder ambient mantle condition and suggests rigid plate motion towards the end of the Mesoarchean if not earlier.

[1] Herzberg, C., and Asimow, P. D. (2015), PRIMELT3 MEGA.XLSM software for primary magma calculation: Peridotite primary magma MgO contents from the liquidus to the solidus, *Geochem. Geophys. Geosyst.*, 16, 563– 578, doi:10.1002/2014GC005631.

[2] Ganne, Jérôme & Feng, Xiaojun. (2017). Primary magmas and mantle temperatures through time. *Geochemistry, Geophysics, Geosystems*. 18. n/a-n/a. 10.1002/2016GC006787.

[3] Singh, S.P., Subramanyam, K.S.V., Manikyamba, C., Santosh, M., Rajanikanta Singh, M., Chandan Kumar, B., 2018. Geochemical systematics of the Mauranipur-Babina greenstone belt, Bundelkhand Craton, Central India: Insights on Neoproterozoic mantle plume-arc accretion and crustal evolution. *Geosci. Front.*

[4] Aulbach, S., Arndt, N.T., 2019. Eclogites as palaeodynamic archives: Evidence for warm (not hot) and depleted (but heterogeneous) Archaean ambient mantle. *Earth Planet. Sci. Lett.* 505, 162–172. <https://doi.org/10.1016/j.epsl.2018.10.025>