## Exploring the use of <sup>17</sup>O-excess in CO<sub>2</sub> for estimating mesophyll conductance of C<sub>3</sub> and C<sub>4</sub> plants

**THOMAS RÖCKMANN<sup>1</sup>**, GETACHEW AGMUAS ADNEW<sup>1</sup>, THIJS PONS<sup>2</sup>, GERBRAND KOREN<sup>2</sup> AND WOUTER PETERS<sup>3</sup>

<sup>1</sup>Institute for Marine and Atmospheric Research Utrecht, Utrecht University

<sup>2</sup>Utrecht University

<sup>3</sup>Wageningen University

Presenting Author: t.roeckmann@uu.nl

Mesophyll conductance to  $CO_2(g_m)$  is an important parameter controlling plant photosynthesis and water use efficiency, and vegetation-atmosphere CO<sub>2</sub> exchange. Here we demonstrate the potential to estimate  $g_m$  from measurements of  $\delta^{17}O$  of  $CO_2$ , where  $\delta^{17}O$  quantifies the <sup>17</sup>O-excess compared to what is expected from the <sup>18</sup>O content of CO<sub>2</sub> according to massdependent isotope fractionation. The  $g_m$  calculations are applied to measurements of  $\delta^{18}$ O and  $\delta^{17}$ O in leaf cuvette gas exchange measurement with sunflower, ivy, and maize, using normal and slightly <sup>17</sup>O-enriched CO<sub>2</sub>. We show that  $g_{m\delta 17}$  estimates can complement and potentially improve the  $g_{m18}$  estimates in settings where the  $\delta^{18}$ O leaf water varies strongly throughout the day, affecting the  $\delta^{18}$ O (CO<sub>2</sub>) difference between the intercellular air space and the CO<sub>2</sub>-H<sub>2</sub>O exchange site. This is because  $\delta^{17}$ O is less sensitive to the unknown isotope fractionation during evapotranspiration of leaf water than  $\delta^{18}$ O because the isotope fractionation processes involved are all mass-dependent. The main limitation to the use of  $g_{m\delta 17}$  is the uncertainty in the measurement of the very small signals in  $\delta^{17}$ O. In general, the precision of a  $g_m$  determination with oxygen isotope techniques decreases when the isotopic difference between CO2 in the intercellular air space and at the CO2-H2O oxygen exchange site becomes very small. In leaf cuvette experiments this limitation can partially be overcome by using <sup>17</sup>O- or <sup>18</sup>O-enriched CO<sub>2</sub>. However, both  $g_{m18}$  and  $g_{m\delta17}$  are not good tracers for plant species with very high mesophyll conductance, because in this case the oxygen isotope gradient between the CO<sub>2</sub>-H<sub>2</sub>O exchange site and the intercellular air space is very small. Both  $g_m$ estimates are sensitive to the assigned degree of equilibration between CO<sub>2</sub> and water ( $\theta_{eq}$ ).