

# Experimental use of olivine-hosted melt inclusions to explore the thermodynamics and kinetics of Fe and S redox equilibria in silicate melts

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Natural silicate melts contain sulfur (dissolved as S<sup>2-</sup> and S<sup>6+</sup>), an important participant in redox chemistry. However, low-pressure experiments used to calibrate Fe oxybarometers, which relate melt Fe<sup>3+</sup>/Fe<sup>2+</sup> to oxygen fugacity ( $fO_2$ ), are based on sulfur-free compositions and this leads to uncertainties in calculating the  $fO_2$  of S-bearing glasses based on their measured Fe<sup>3+</sup>/Fe<sup>2+</sup>[1]. Furthermore, if electron exchange between dissolved S and Fe is temperature-dependent, then liquid S<sup>6+</sup>/S<sup>2-</sup> and Fe<sup>3+</sup>/Fe<sup>2+</sup> ratios will change during cooling. Olivine-hosted melt inclusions (MIs) equilibrate rapidly with the  $fH_2$  and  $fH_2O$  in 1-atm furnaces[2] while retaining their sulfur. This offers a way to explore electron exchange between S and Fe in silicate melts as a function of  $T$  and  $fO_2$ , and to assess whether melts equilibrated at the same  $T$  and  $fO_2$  but cooled at different rates and quenched at different temperatures record the same S and Fe speciation.

Olivine-hosted MIs were held at 1225°C for 24–48 hours at nine  $fO_2$ s ( $\Delta FMQ-3$  to  $\Delta FMQ+3$ ) and then drop-quenched ( $>700^\circ C/s$ )[3] or cooled at rates as low as  $\sim 30^\circ C/s$  and drop-quenched at lower  $T$ . A subset of the glasses was analyzed using S and Fe  $\mu$ -XANES. Based on furnace  $fO_2$ s, the measured Fe<sup>3+</sup>/Fe<sup>2+</sup> in the glasses are low relative to estimates from published Fe oxybarometers[1]. The S<sup>6+</sup>/S<sup>2-</sup> of the MI glasses in the 1225°C experiments follow the sigmoidal dependence on  $fO_2$  determined for hydrous basalts[4]. At the same  $fO_2$ , the cooling-rate experiments have systematically lower S<sup>6+</sup> than MIs quenched from 1225°C. Two unheated MIs from Mauna Loa with cooling rates[3] of  $\sim 1^\circ C/s$  and  $\sim 130^\circ C/s$  follow the same sigmoidal relationship between  $\Delta FMQ$  and S<sup>6+</sup>/S<sup>2+</sup> as the experimental inclusions. The  $\sim 1^\circ C/s$  natural MI retained  $\sim 90\%$  of S as S<sup>6+</sup> as it cooled to  $T < 900^\circ C$ [3], consistent with a relatively weak  $T$ -dependence for S-Fe redox equilibria in Hawaiian tholeiitic melts. However, the kinetics of S-Fe electron exchange are sufficiently rapid to modify S speciation in melts cooled at rates lower than  $\sim 30^\circ C/s$ .

[1] Graz *et al.* (2006) *AGU*, V43B-01; [2] Gaetani *et al.* (2012) *Geology* 40, 915; [3] Saper & Stolper (2020) *G<sup>3</sup>* 21, e2019GC008772; [4] Jugo *et al.* (2010) *GCA* 74, 5926