

An Observationally Constrained, ^{234}Th -Derived Global POC Flux Model

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The biological carbon pump (BCP) performs a significant role in mitigating increasing anthropogenic CO_2 emissions and consequent heating effects [1]. Quantifying the contribution of the BCP to carbon sequestration has been the topic of several recent studies, with the average yearly POC flux out of the upper ocean ranging from 4 to 12 PgC year^{-1} [2].

Here we use the commonly employed Thorium-234 (^{234}Th) proxy to determine carbon fluxes throughout the world's oceans. The disequilibrium between ^{234}Th (half-life ~ 24.1 days), a highly particle-reactive radionuclide, with its parent, Uranium-238 (^{238}U , half-life $\sim 4.5 \times 10^9$ years), can be used to quantify ^{234}Th fluxes out of the surface ocean [3]. Along with the particulate ratio of ^{234}Th with POC ($^{234}\text{Th}:\text{POC}$), these measurements have been used to estimate POC fluxes throughout the water column [4].

Using a global database, over 60,000 global measurements of ^{234}Th , ^{238}U , and $^{234}\text{Th}:\text{POC}$ were regressed to a 2.8-degree grid using a suite of Machine Learning (ML) and Minimum Variance algorithms [5]. This gridded data was then used to drive a 3D global model of POC fluxes out of the regionally varying Euphotic Zone (EZ) using a set of sparse, implicit and explicit transport matrices derived from a global, coupled General Circulation Model [6].

Our model captures known major features of ^{234}Th , with an average of 2.03 ± 0.20 dpm L^{-1} globally and greatest activities in ocean gyres. The model estimates an average of ~ 6.3 PgC year^{-1} of POC out of the EZ, with an average flux of 44 ± 43 $\text{mgC m}^{-2} \text{day}^{-1}$ and greatest fluxes in the coastal, equatorial, and extreme latitude regions. Furthermore, our model predicts that the Pacific Ocean contributes to $\sim 50\%$ of the annual flux, with ~ 3 PgC year^{-1} , followed by the Atlantic Ocean, with a $\sim 30\%$ contribution of ~ 2.1 PgC year^{-1} .

[1] Ducklow et al. (2001), *Oceanography* 14, 50–58

[2] DeVries & Weber (2017), *Global Biogeochem. Cycles* 31, 535–555

[3] Buesseler et al. (1995) *Deep Sea Res. PII* 42, 777–804

[4] Buesseler et al. (2006), *Marine Chemistry* 100, 213–233

[5] Ceballos-Romero et al. (In prep.)

[6] Khatiwala et al. (2005) *Ocean Model.* 9, 51–69

