

## **Thermodynamic and kinetic constraints on molecular hydrogen abundance and isotope systematics in hydrothermal fluids**

**DR. ANDREA RICCI, PHD<sup>1</sup>**, JENS FIEBIG<sup>2</sup>, BARBARA I. KLEINE<sup>1</sup>, JÓHANN GUNNARSSON-ROBIN<sup>1</sup>, KENNEDY M. KAMUNYA<sup>3</sup>, BRUCE MOUNTAIN<sup>4</sup> AND ANDRI STEFÁNSSON<sup>1</sup>

<sup>1</sup>University of Iceland

<sup>2</sup>Goethe University Frankfurt

<sup>3</sup>Kenya Electricity Generating Company Ltd.

<sup>4</sup>GNS Science

Presenting Author: [andrearicci@hi.is](mailto:andrearicci@hi.is)

Molecular hydrogen is one of the major gas constituents in hydrothermal fluids where it plays a pivotal role in geological and biological processes. Nonetheless, factors controlling molecular hydrogen abundance and isotope fractionation remain poorly understood and quantified. We report hydrogen fugacity ( $f_{H_2}$ ) and  $\delta D$  of  $H_2$  and  $H_2O$  for hydrothermal fluids of variable temperature (226-359 °C) from terrestrial volcanic arc and rift systems sourced by seawater and meteoric water. The hydrothermal fluid  $f_{H_2}$ ,  $\delta D-H_2$  and  $\delta D-H_2O$  values are of 0.002-3.3 bar, -646 to -391‰ and -94.1 to +11.3‰, respectively. Comparison of dataset with results of geochemical modeling revealed that in meteoric water systems  $H_2$  production is controlled by the reduction of  $H_2O$  upon oxidation of aqueous  $Fe^{II}$  to  $Fe^{III}$  and subsequent formation of Fe-containing mineral. Elevated sulfur contents, sourced from volcanic gas and/or seawater, result in oxidation of aqueous  $S^{II}$  to pyrite and  $S^{+VI}$ , providing a further source of electrons along with  $Fe^{II}$  oxidation. Our results show that hydrogen fugacity in hydrothermal fluids is controlled by metastable equilibria along a fluid-rock reaction path that primarily depend on temperature, rock-to-water ratio, source water composition and volcanic gas input. The dataset demonstrates that  $\delta D$  ratio of  $H_2$  is controlled by the isotopic composition of the source water and equilibrium isotope fractionation at the hydrothermal reservoir temperatures. Upon fluid ascent to surface,  $H_2-H_2O$  isotope exchange re-equilibration may occur, this depends on cooling rate coupled with the kinetics of the isotope fractionation reaction. Based on the extent of hydrogen isotope disequilibrium, we estimate reservoir-to-surface travel times of minutes to <3 hours and few hours to days for geothermal wells and fumaroles fluids, respectively. The Project has received funding from the European Union's Horizon 2020 under Grant Agreement #818169.