

Isotope signs ($^{234}\text{U}/^{238}\text{U}$, ^2H , ^{18}O) for the paleo-permafrost existence in European Russia

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Global climatic variations have a significant impact on isotopic and chemical composition of the underground hydrosphere. In particular, glacial meltwater formed from the permafrost degradation, as a rule, has a light composition of stable isotopes ($\delta^{13}\text{C}$, $\delta^{18}\text{O}$, and $\delta^2\text{H}$) [1]. Apparently, this water should also be marked with excess uranium-234 [2]. We studied groundwater in the central part of the East European Plain on the right bank of the Volga River (middle course) to a depth of 400 m. During the Last Glacial Period, there was an area of continuous permafrost with a thickness of up to 200–250 m and with temperatures of about -5°C [3]. Comparison isotopic and chemical data shows the three-component mixture of groundwater. The first component is modern water, which is unambiguously recorded by the presence of tritium (^3H). It is freshwater with salinity $S < 0.5$ g/L, it has the least depletion in the isotopic composition of hydrogen and oxygen with $\delta^{18}\text{O} \rightarrow -12.9\text{‰}$ and $\delta^2\text{H} \rightarrow -90\text{‰}$ that is close to modern precipitation of the region, and practically equilibrium uranium $^{234}\text{U}/^{238}\text{U} \rightarrow 1$. The second component, more likely, is the water of late and postglacial genesis. It is isotopically lightest water with $\delta^{18}\text{O} \rightarrow -17.0\text{‰}$ and $\delta^2\text{H} \rightarrow -119\text{‰}$ (maybe it is the water of glacial lake?), it has a slight increase in salinity $S \approx 0.5\text{--}1.5$ g/L, and a small excess of uranium-234 $^{234}\text{U}/^{238}\text{U} \approx 4$. The third component, as we explained, is the meltwater of the former permafrost. It is brackish water $S \rightarrow 3$ g/L in which a significant growth of salinity is mostly associated with increase in the sulfate content. It could be the result of both presence of evaporites in the water-bearing sediments and the increased solubility of sulfates at near-zero temperature. This water has an intermediate composition of the stable isotopes with $\delta^{18}\text{O} \approx -15.0\text{‰}$ and $\delta^2\text{H} \approx -110\text{‰}$, and maximum disequilibrium in uranium $^{234}\text{U}/^{238}\text{U} \rightarrow 15.7$. The abnormal excess of uranium-234 in groundwater is probably the result of its predominantly leaching from a mineral lattice of water-bearing sediments during the thawing of permafrost.

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[1] Vasil'chuk et al.(2018)Geoscience Frontiers,2,471-483 [2] Tokarev et al.(2021)Geochemistry Int. in press. [3] Sychova(2012)Cryosphere,4,45-56.