## Volatiles in melt inclusions and apatite from Erebus phonolite

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Magmatic volatiles such as CO<sub>2</sub> and H<sub>2</sub>O play a major role in magmatic processes and eruptive behaviour. Abundances of H<sub>2</sub>O in the melt can be determined by volatile measurements on apatite, combined with thermodynamic analyses of F-Cl-OH partitioning between apatite and melt [1]. A similar method should be possible to assess melt CO<sub>2</sub> contents but studies of the concentration of CO<sub>2</sub> in apatite and its partition behaviour are few. Here we provide new measurements (using SIMS) and thermodynamic analysis of CO<sub>2</sub>, H<sub>2</sub>O, F, Cl and S concentrations in co-existing apatite and phonolitic glass inclusions, hosted by anorthoclase megacrystals of lava bomb samples from Mt. Erebus (Antarctica; [2]). We find that the melt inclusions (MIs) contain similar volatile compositions: ~0.16 wt.% H<sub>2</sub>O, ~90 ppm CO<sub>2</sub>, ~2100 ppm F, ~1400 ppm Cl, ~350 ppm S. The coexisting fluorapatite crystals have a limited range of  $F \approx$ 3.5 wt.%, Cl  $\approx$  0.13 wt.%, H<sub>2</sub>O  $\approx$  0.06 wt.%, and S  $\leq$  5 ppm, but a wide variation in CO<sub>2</sub> from  $\approx$  120 to 1600 ppm. Apatite-melt inclusion pairs indicate exchange coefficient  $(K_{p})$  for Cl-F of ~0.11, equivalent to that calculated from [1] at 950–1000 °C. Using the apatite hygrometer model of [1], the melt should contain >0.5 wt.% H<sub>2</sub>O. This is higher than that measured in MIs but can be explained by diffusive H<sub>2</sub>O loss. The calculated values of the Nernst partition coefficient (D) for  $CO_2$  are ~1.6–2.8, which are much higher than those (of ~0.11–0.12) determined by [3] at 1 GPa and 1250 °C for trachytic melts and F-rich apatite. The calculated value of exchange coefficient K<sub>D</sub> for H<sub>2</sub>O-CO<sub>2</sub> (on the basis of wt.%) is ~0.04 (±0.02), ~10 times lower than that determined at 1250 °C by [3]. This discrepancy can be explained by a temperature difference of ~300 °C, and is comparable to the difference in K<sub>p</sub> for OH-F between 950 and 1250 °C [1]. More experimental studies on apatite-melt CO<sub>2</sub> partitioning are needed to improve estimates of melt CO<sub>2</sub> budgets from CO<sub>2</sub> in apatite.

[1] Li & Costa (2020) *GCA* **269**, 203-222. [2] Moussallam *et al.* (2015) *EPSL* **413**, 1-12. [3] Riker *et al.* (2018) *Am Mineral* **103**, 260-270.