

Does the polymorphism control the boron isotopic composition of biocarbonate?

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The past variations of the atmospheric partial pressure of CO₂ can be constrained either from direct measurements in air bubbles trapped in ice cores or by geochemical proxies from oceanic sedimentary records. Among these geochemical proxies, the boron isotopic composition ($\delta^{11}\text{B}$) of carbonates is one of the most used. Indeed, carbonates are supposed to incorporate borate ions of the solution from which they precipitate, and thus their $\delta^{11}\text{B}$ depends on solution pH (and so on dissolved atmospheric CO₂) (eg Vengosh et al, 1991; Hemming and Hanson, 1992). However, Noireaux et al (2015) showed that the crystallography of the calcium carbonate inorganically precipitated could also control the boron isotopic composition; aragonite $\delta^{11}\text{B}$ values reflect the $\delta^{11}\text{B}$ of solution borate, whereas calcite $\delta^{11}\text{B}$ values are enriched in ¹¹B compared to the $\delta^{11}\text{B}$ of solution borate, likely because of the partial incorporation of boric acid. Intriguingly, the published data on biogenic carbonates (corals, foraminifera, brachiopods...) show the opposite behaviour with a larger boron isotopic fractionation relative to seawater borate ion for bio-aragonite compared to bio-calcite. To see if this different behaviour between inorganic and biogenic carbonates is linked to the crystallography and/or to biological processes, we will measure the $\delta^{11}\text{B}$ and B/Ca ratios of various shells presenting both aragonitic and calcitic layers (*Mytilus edulis*, *Haliotis rufescens*, *H. laevigata*, *Patella* sp., *Nerita* sp.). The first results show a clear difference in boron content, the aragonitic layer being 3 to 10 times enriched in boron compared to the calcitic layer, as expected from inorganic precipitation experiments (Mavromatis et al, 2015).

Hemming and Hanson (1992), GCA, 56, 537-543.

Mavromatis et al (2015), GCA, 150, 299-313.

Noireaux et al (2015) EPSL, 430, 398-407.

Vengosh et al (1991), GCA, 55, 2901-2910.