Vanadium isotopic composition of garnet-clinopyroxene pairs from kimberlite-hosted mantle xenoliths

PAGE, S. E., ¹ Prytulak, J., ¹ Nowell, G^1 ., Novella, D^2 ., Yaxley, G. M., ³

¹Dept. of Earth Sciences, Durham University, DH1 3LE, UK (correspondance: sophie.e.page@durham.ac.uk)

²Dept. of Earth Science, Cambridge University, CB2 3EQ, UK

³Research School of Earth Science, The Australian National University, Canberra, ACT, 2601, Australia

Prevailing mantle oxygen fugacity (fO_2) is hotly debated, since current fO_2 proxies have large uncertainties of at least half an order of magnitude. Thus, new approaches to quantify fO_2 are sought.

Vanadium (V) is a moderately incompatible, immobile transition metal with multiple valence states (+2, +3, +4, +5). Previous studies have shown that V concentrations in magmatic rocks can respond to variable fO_2 [1], however this approach is hindered by uncertainties in mantle source concentrations and degree of melting. Variations in the stable isotope compositions of multi-valent elements may also respond to fO_2 . Theoretically, heavier stable isotopes preferentially enter sites with stiffer bonding environments (i.e. lower coordination and higher valence states). Hence, V stable isotopes could be used as a redox proxy, but the effects of coordination and oxidation state must be evaluated.

Here we present V stable isotope data for natural mineral separates. Garnet and clinopyroxene in chemical and textural equilibrium were picked from fresh, well-characterised, lherzolitic, mantle xenoliths from the Diavik diamond mine in the central Slave craton, Canada [2]. Co-ordination of V in both garnet and clinopyroxene is VI-fold; therefore, valence state should be the primary variable affecting bond strength. Determined bulk fO_2 ranges from -2.38 $\Delta \log fO_2^{[FMQ]}$ to -3.36 $\Delta \log fO_2^{[FMQ]}$ [2]. The measured garnet-clinopyroxene intermineral fractionation factor is negligible. Neither garnet nor clinopyroxene δ^{51} V in this suite co-vary with garnet Fe³⁺/ Σ Fe or bulk fO₂. However, bulk rock δ^{51} V ranges from -1.08 to -1.73, positively correlating with fO_2 , suggesting an additional redox-sensitive phase. The Diavik xenoliths are kimberlitehosted, and given the significant difference in δ^{51} V between the main mineral hosts of V and bulk powders, the effect of host contamination needs further examination. This variance underlies the ambiguity inherent to using rock powders without mineralogical context.

[1] Canil, D. (1997) *Nature*. **389**, p. 842-845 [2] Yaxley et al. (2017) *Sci. Reports*. **7** (30)