

Li-isotopic evidence for fluid exsolution and remobilisation of REE-HFSE in peralkaline Siwana granite

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The peralkaline Siwana granite of the Neoproterozoic Malani Igneous Suite in western India hosts a diverse suite of rare earth elements (REE) and high field strength elements (HFSE)-bearing minerals formed during magmatic to post magmatic evolution of the granite body. Extensive fluid-induced deuteric to hydrothermal alteration led to the partial breakdown and re-equilibration of primary igneous minerals and remobilization and redistribution of precursor REE and HFSE minerals.

In this study, Li isotopes were measured in-situ on magmatic and deuteric/hydrothermal zones of coexisting clinopyroxene and amphibole using LA-MC-ICPMS. The early magmatic hedenbergite furnish $\delta^7\text{Li}$ of $20.4 \pm 3.2\%$ while the late-magmatic aegirine have $\delta^7\text{Li}$ of 8.6 ± 2.6 to $-16.7 \pm 0.65\%$. Hydrothermal aegirine rims have heavier $\delta^7\text{Li}$ of $24.6 \pm 0.8\%$. Amphiboles display a similar Li-isotopic trend with early magmatic richterite having $\delta^7\text{Li}$ of $18.5 \pm 0.65\%$ and late magmatic riebeckite-arfvedsonite having lighter $\delta^7\text{Li}$ of $-3.5 \pm 1.4\%$. Hydrothermal arfvedsonites replacing the magmatic amphiboles have much heavier $\delta^7\text{Li}$ of $29 \pm 0.8\%$ to $33.3 \pm 0.7\%$. The trend of decreasing $\delta^7\text{Li}$ in the magmatic pyroxenes and amphiboles can be explained to be due to progressive degassing and exsolution of ^7Li -rich fluids from the peralkaline magmas during low-pressure vapour-saturated crystallization [1]. The hydrothermal minerals crystallized from the exsolved fluids which explains their heavier Li isotopic compositions. The hydrothermal amphiboles are characterized by high F, Cl, and Pb. The Cl-rich exsolved fluids were responsible for the mobilization of the REE and HFSE as chloride complexes while F assisted the re-precipitation of hydrothermal REE minerals [2].

[1] Vlastélic et al. (2011) *Chem. Geol.* **284**, 26–34 [2] Migdisov et al. (2016) *Chem. Geol.* **439**, 13–42.