## Carbonate Fluorescence and Carbonate Stable Isotopes Footprint around Ferrobamba Cu-Skarn, Southern Peru

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Skarns in the Las Bambas district of southern Peru, are hosted by marine sedimentary sequences deposited during the Cretaceous. These sedimentary rocks were intruded by the Andahuaylas-Yauri batholith in the Eocene-Oligocene [1]. A series of plutons, stocks and dikes emplaced into limestones generated skarn alteration and copper mineralization at three main porphyry-skarn mineralization centres: Ferrobamba (1,351 Mt @ 0.62%Cu), Chalcobamba (315 Mt @ 0.57%Cu) and Sulfobamba (184Mt @ 0.62%Cu) [2].

Calcite veins around Ferrobamba have three different shortwave ultra-violet fluorescent colours: blue, pink, and bright white. Carbon and oxygen isotope analyses have shown that the fluorescent response of calcite varies systematically with its isotopic compositions.

Blue fluorescent calcite has high  $\delta^{13}$ C and high  $\delta^{18}$ O, similar to the isotopically unaltered limestones and marbles. Pink fluorescent calcite has low  $\delta^{13}$ C and low  $\delta^{18}$ O, consistent with carbonate veins derived from magmatic-hydrothermal fluids. A continuous array of carbonate isotope composition from depleted C<sup>13</sup> and O<sup>18</sup> at the centre of Ferrobamba, to enriched C<sup>13</sup> and O<sup>18</sup> on the margins of the deposit, defines an isotopic halo 5 x 3 km in size around Ferrobamba.

This broad zoned pattern is disrupted by bright white fluorescent calcite, which has anomalous low  $\delta^{13}C$  and high  $\delta^{18}O$  values. This variety of calcite is related to late breccias with low-temperature smectite alteration.

The high  $\delta^{13}$ C-high  $\delta^{18}$ O to low  $\delta^{13}$ C-low  $\delta^{18}$ O trend is the result of the interaction between magmatic water and the limestones. Whereas, the low  $\delta^{13}$ C-high  $\delta^{18}$ O signature is the product of a cooling magmatic-hydrothermal fluid that didn't undergo water-rock interaction.

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