

Copper mobilisation and isotope fractionation during subduction

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Subduction zones are key components of global geochemical cycles. Fluids released from subducted slabs generate arc magmatism and carry an elemental budget into the mantle wedge that is consequently not subducted into the deeper mantle alongside the residual slabs, thus affecting the surface to mantle mass balance. The fluids also appear to be oxidising [1]. Yet how and where the oxidising properties are attained has so far been largely speculative. Potential oxidising agents transported by slab-derived fluids include sulfate, carbonate, and Fe³⁺ [e.g. 1,2].

The dehydration of serpentinites underneath the subducting mafic oceanic crust has been suggested as a major source of H₂O in subduction zones [3]. We explore whether the oxidizing properties of slab fluids are related to the dehydration of serpentinites or generated at a later stage as the fluids pass through the mafic crust, subducted sediments, and the sub-arc mantle. For this purpose, we focus on Cu isotope ratios and chalcophile element concentrations in serpentinites from the Western Alps [4]. Our samples represent various stages of prograde metamorphic reaction and slab devolatilisation from the replacement of lizardite by antigorite to the first stages of serpentinite dehydration at eclogite facies.

Copper concentrations in the serpentinites decrease by about 50 % from ~16 ppm to ~8 ppm with increasing degree of metamorphism, indicating that Cu is affected by open system processes during devolatilisation. The loss of Cu from the slab is consistent with earlier studies suggesting the mobilisation of Fe and Zn with associated isotope fractionation [5,6] and highlight the significant mass transfer of compatible and chalcophile elements from subducted serpentinites during devolatilisation.

[1] Kelley and Cottrell (2009) *Science* 325, 605-607, [2] Evans et al. (2012) *Geology* 40, 783-786, [3] Ulmer and Trummsdorff (1995) *Science* 268, 858-861 [4] Debret et al. (2013) *Chemical Geology* 357, 117-133 [5] Debret et al. (2016)

Geology 44, 215-218 [6] Pons et al. (2017) Nature
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