

Origins of a global marine late Silurian C isotope excursion: Insights from stable metal isotopes (Ca, Sr, Cr) from Gotland and Prague Basin

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The causes and *primary* versus *diagenetic* origins of the globally-recognized large positive C isotope excursions (CIE), recorded in the Paleozoic marine carbonates, and their purported links to coeval changes in the global C cycle are controversial and highly disputed [1, 2]. The largest positive $\delta^{13}\text{C}$ excursion of the entire Phanerozoic, i.e., the mid-Ludfordian CIE (~9‰), is documented in the Late Silurian marine carbonates worldwide [3]. Recently, it has been proposed that the origin of this CIE is related to a purported 'carbonate hypersaturation' of the late Silurian surface oceans [4], and the associated kinetically-controlled evasion of isotopically light CO_2 (and CH_4) gases from the surface oceans, coeval also with a rapid carbonate precipitation [4], and rate-controlled effects documented by Ca isotopes [5]. An alternative explanation suggest that this CIE and the above rate-controlled effects are rather due to an early marine diagenesis and post-depositional equilibration of primary marine carbonates with paleo-seawater or seawater-derived fluids, leading to resetting of primary isotope signals [2]. To further test these two plausible scenarios (i.e., a global seawater *hypersaturation* vs. *early marine diagenesis*), we present here stable Ca, Sr and Cr isotope records of Late Silurian marine carbonates from two remote paleo-locations, represented by Gotland (Sweden) and Prague Basin (Czech Rep.). Importantly, both sites yielded generally coherent $\delta^{44/40}\text{Ca}$, $\delta^{88/86}\text{Sr}$ and $\delta^{53}\text{Cr}$ trends across the CIE, corroborating the global nature of these isotope trends. We will discuss the implications of these results for the *primary* versus *diagenetic* origins of the late Silurian CIE, and also for the reconstructions of paleo-seawater metal isotope signatures.

[1] Saltzman & Edwards (2017) *EPSL*, **464**, 46–54. [2] Ahm et al. (2018) *GCA*, in press. [3] Fryda & Manda (2013), *Bull. Geosci.* **88**, 463–482. [4] Kozłowski (2015) *Bull. Geosci.* **90**, 807–840. [5] Farkas et al. (2016) *EPSL*, **451**, 31–40.