

## Heavy $\delta^{30}\text{Si}$ in Archean granitoids as evidence for supracrustal components in their sources

LUC ANDRÉ<sup>1</sup>, KATHRIN ABRAHAM<sup>2</sup>, STEPHEN F. FOLEY<sup>3</sup>, AND AXEL HOFMANN<sup>4</sup>

<sup>1</sup> Earth Sciences Department, Royal Museum for Central Africa, Leuvensesteenweg 13, Tervuren, Belgium.

<sup>2</sup> University of Oxford, Department of Earth Sciences, Oxford, United Kingdom

<sup>3</sup> Dept. Earth and Planetary Sciences, Macquarie University, North Ryde 2109, Australia.

<sup>4</sup> Department of Geology, University of Johannesburg, South Africa.

We report  $\delta^{30}\text{Si}$  data from 37 samples of 17 Archean (3.5-2.7Ga) felsic intrusive rocks from the Kaapvaal craton. They are representative of three types of plutons: (a) Trondhjemite-Tonalite-Granodiorite (TTG) (n=13); (b) Granite-Monzogranite-syenogranite (GMS) (n=15); (c) composite TTG-GMS (n=9). Two to 4 aliquots of each specimen powder were prepared to control the full chemical reproducibility. Each aliquot solution was then analysed 7 to 12 times on two MC-ICP-MS instruments (Nu Plasma II and Thermo Neptune) with an average reproducibility ( $2\sigma_M$ ) of 0.05‰. In terms of their average  $\delta^{30}\text{Si}$ , we find no resolvable isotopic differences between TTGs, TTG-GMSs and GSMs [ $\delta^{30}\text{Si} = -0.01 \pm 0.07\text{‰}$  ( $\pm 2\sigma_X$ ),  $-0.02 \pm 0.09\text{‰}$ ,  $-0.02 \pm 0.13\text{‰}$ , respectively], but they all contrast significantly with coevally run data for the GA granite standard ( $\delta^{30}\text{Si} = -0.22 \pm 0.08\text{‰}$ ). The grand average for these Archean plutons ( $\delta^{30}\text{Si} = -0.02 \pm 0.02\text{‰}$ ,  $\pm 2\sigma_M$  for n=37) is significantly heavier than Phanerozoic I- and A-type granites ( $\delta^{30}\text{Si} = -0.19 \pm 0.02\text{‰}$ ,  $\pm 2\sigma_M$  for n=27 [1,2]), and rhyolitic liquids differentiated from basalts ( $\delta^{30}\text{Si} = -0.19 \pm 0.02\text{‰}$  [3]). We do not see any significant  $\delta^{30}\text{Si}$  trends with pluton age and  $\text{SiO}_2$  contents, nor with the large variations in  $\text{La}_N/\text{Yb}_N$  (5 to 68) and  $\text{Sr}/\text{Y}$  (0.25 to 250). Therefore, melting or crystallisation in the garnet or plagioclase stability fields can be ruled out as a cause for this heavy signature, the origin of which must be rooted in the source protoliths. Our preferred explanation is the melting of recycled silicified basalts ( $\delta^{30}\text{Si} = +0.49 \pm 0.54\text{‰}$  [4]) and cherts ( $\delta^{30}\text{Si} = +0.60 \pm 0.44\text{‰}$  [4]) which are common supracrustal rocks during the Eo- and Paleo- Archean Era.

[1] Savage et al. (2012) GCA 92, 184-202. [2] Poitrasson & Zambardi (2015) GCA 167, 301-312. [3] Savage et al. (2011) GCA 75, 6124-6139 [4] Abraham et al. (2011) EPSL 301, 222-230.