

## Mass-dependent isotopic fractionation of argon in natural systems: Implications to $^{40}\text{Ar}/^{39}\text{Ar}$ dating

B.D. TURRING<sup>1\*</sup>, D.E. CHAMPION<sup>2</sup>, AND C.C. SWISHER<sup>1</sup>  
III<sup>1</sup>

<sup>1</sup>Department of Earth and Planetary Sciences Rutgers, The  
State University of New Jersey, Piscataway, NJ 08854  
USA (\*correspondence: [bturrin@eps.rutgers.edu](mailto:bturrin@eps.rutgers.edu),  
[cswish@eps.rutgers.edu](mailto:cswish@eps.rutgers.edu))

<sup>2</sup>U.S. Geological Survey, 345 Middlefield Rd. Menlo Park  
CA 94025 ([dchamp@usgs.gov](mailto:dchamp@usgs.gov))

Like all radiometric dating methods, in order to obtain accurate ages from the K-Ar radioisotope chronometer system, an accurate assessment and correction of the initial concentration of the daughter isotope (often referred to as either the “initial” or “trapped” component)  $^{40}\text{Ar}$  is required. Generally it is assumed that the total amount of  $^{40}\text{Ar}$  is composed of radiogenic  $^{40}\text{Ar}^*$  derived from the decay of  $^{40}\text{K}$  and a trapped  $^{40}\text{Ar}_{\text{trp}}$  component derived from the atmosphere, incorporated into the system in proportion the atmospheric  $^{40}\text{Ar}/^{36}\text{Ar}$  ratio 298.56 [1,2]. If this assumption is false, then the resultant calculated ages are incorrect. For the most part this assumption seems to hold true. However, evidence of non-atmospheric  $^{40}\text{Ar}/^{36}\text{Ar}$  isotopic ratios that are enriched in  $^{36}\text{Ar}$  are common in young volcanic rocks [3,4,5,6,7,8]. In  $^{38}\text{Ar}/^{36}\text{Ar}$  vs  $^{40}\text{Ar}/^{36}\text{Ar}$  space these  $^{36}\text{Ar}$  samples lie on the kinetic mass dependent fractionation trend. Results from analyses reported here support Renne [9] that “kinetic fractionation of Ar isotopes is common, if not ubiquitous, in volcanic rocks”.

To address this issue, we present an analytical protocol that identifies, and corrects for kinetic mass dependent fractionation that can be applied to the  $^{40}\text{Ar}/^{39}\text{Ar}$  dating method. Our protocol further accounts for temporal variations in the Terrestrial atmospheric  $^{40}\text{Ar}/^{36}\text{Ar}$  ratio [10,11,12].

[1] Lee et al. (2006) *Geochim Cosmochim Acta* **70**, 4507-4512. [2] Valkiers et al. (2010) *Intl. Jour. Mass Spectrometry* **291**, 41-47. [3] Dalrymple (1969) *EPSL* **6**, 47-55. [4] Krummenacher (1970) *EPSL* **8**, 109-117. [5] Kaneoka (1980) *EPSL* **48**, 284-292. [6] Masumoto et al. (1989) *Mass Spectroscopy* **37**, 353-363. [7] Masumoto & Kobayashi (1995) *Chem. Geology* **125**, 123-135. [8] Ozawa et al. (2006) *Rev. Min. & Geochem* **47**, 411-480. Renne et al. (2009) *Quat. Geochron.* **4** 288-298. Cadogan PH (1977) *Nature* **268** 38-41 Bender et al. (2008) *Proc. Natl. Acad. Sci.* **105**, 8232-8237. Pujol et al. (2013) *Nature* **498**, 97-90