

Large $^{238}\text{U}/^{235}\text{U}$ Fractionation in an alkaline, euxinic lake: implications for the marine sedimentary record

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Variations in the ratio of $^{238}\text{U}/^{235}\text{U}$ in sedimentary rocks have been used to quantify the oxygenation of the Precambrian oceans and changes in the extent of ocean anoxia at the P-T boundary. Despite research demonstrating that U(VI) reduction reactions are the primary driver of fractionation ($\epsilon^{238}\text{U} \sim -1\%$), interpretations of $^{238}\text{U}/^{235}\text{U}$ remain ambiguous. A recent review of modern marine sediments suggests that the expression of this fractionating U reduction in all anoxic and hypoxic basins is limited by diffusion of aqueous uranium into the sediment column, limiting observed shifts to $\epsilon^{238}\text{U} \leq 0.6\%$. In contrast, ancient sediments from the Black Sea suggest $\epsilon^{238}\text{U}$ greater than this diffusive limit [1].

To address this disparity, we characterized the U cycle in Mono Lake (eastern California, USA) in order to better understand the physical and chemical controls on $^{238}\text{U}/^{235}\text{U}$ fractionation in anoxic settings. Mono Lake is a closed basin alkaline lake (pH 9.8) with characteristically long residence times for actinides. Evaporation has resulted in highly concentrated uranium (500 $\mu\text{g/L}$) and carbonate (320 mM) within the lake. Anoxic, sulfidic conditions exist at depth in the water column, and support U(VI) reducing conditions.

We characterized the $^{238}\text{U}/^{235}\text{U}$, $^{234}\text{U}/^{238}\text{U}$, and U concentrations of lake water as a function of depth through the water column, as well as inputs (streams and springs) and outputs (carbonate tufas and lake bottom sediments). The lake water $^{238}\text{U}/^{235}\text{U}$ is $-1.1\% \pm 0.1\%$ and is invariant with depth. In contrast, authigenic U in the lake bottom sediments average $0.0\% \pm 0.2\%$ and have constant concentrations over the 60 cm depth interval we measured. Freshwater streams and springs deliver U to the lake with a $^{238}\text{U}/^{235}\text{U} = -0.2\% \pm 0.2\%$. A numerical model of the system suggests that (1) the lake is in steady-state for U and U isotope compositions, (2) U reduction is not limited by diffusion into the sediment and (3) U reduction likely occurs in the water column. These results suggest that inferences of $\epsilon^{238}\text{U} \geq 0.6\%$ in the sedimentary record may indicate the onset of U reduction in the water column

[1] Andersen, *et al.* (2017) *RiMG* **82**, 799-850.