

## **Zinc isotope constraint on the petrogenesis of Fe-rich peridotites and pyroxenites of the Bohemian Massif**

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Zinc isotope compositions of mafic-ultramafic rocks show limited variations during partial melting<sup>[1,2]</sup> but significant diversity during mantle metasomatism<sup>[2,3]</sup>, indicating that Zn isotopes may be a sensitive tracer of metasomatic processes in the mantle and in turn, provide constraints to petrogenesis of mantle-derived rocks. Mantle-derived peridotites and pyroxenites are common in several tectonostratigraphic units of the Bohemian Massif (Czech Republic) and their petrogenesis is a matter of debate<sup>[4,5]</sup>. Here, we present high-precision Zn elemental and isotopic analyses of a suite of Mg-rich and Fe-rich peridotites as well as associated pyroxenites from the Moldanubian Zone of the Bohemian Massif in order to constrain the petrogenesis of these mantle-derived rocks.

Our results show that the Mg-rich peridotites have homogeneous Zn isotopic compositions with  $\delta^{66}\text{Zn}$  of  $0.15 \pm 0.09\%$  (2SD), similar to that of the primitive mantle<sup>[1]</sup>. In contrast, the Fe-rich peridotites have resolvably higher Zn contents and  $\delta^{66}\text{Zn}$  values (0.11 to 0.41‰). The pyroxenites show the largest variations of  $\delta^{66}\text{Zn}$  values from -0.33 to 0.42‰. However, except for three pyroxenites with higher  $\delta^{66}\text{Zn}$  values, the others have Zn contents and  $\delta^{66}\text{Zn}$  values similar to or lower than those of the Mg-rich peridotites. Such compositions of pyroxenites are complementary to those of the Fe-rich peridotites. Thus, our Zn isotopic data indicates that the Fe-rich peridotites were produced by melt-rock interactions between the Mg-rich peridotites and slab-derived basaltic melts with a heavy Zn isotopic signature, while the pyroxenites represent the crystallized products of residual melts<sup>[4]</sup>.

[1] Sossi et al. (2018, CG): 477, 73–84; [2] Huang et al. (2018, JGR, Accepted); [3] Wang et al. (2017, GCA): 198, 151–167; [4] Ackerman et al. (2009, Lithos): 259, 152–167; [5] Svojtka et al. (2016, JP): 57, 897–920.