

## Radiogenic isotope geochemistry of Cabo Ortegal pyroxenites: origin and age of a sub-arc mantle domain

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Pyroxenites exposed in orogenic peridotite massifs illustrate the diversity of petrogenetic processes in supra-subduction environments. However, the source of their parental melts and the timing of their formation remain poorly constrained due to overprinting by metasomatism and metamorphism. We here report Sr-, Nd-, Hf- and Os-isotope compositions of a well-characterized set of pyroxenites from the Variscan Cabo Ortegal Complex, Spain. Previous petrological and geochemical data showed that these pyroxenites formed following melt/rock interaction associated with the intrusion of picritic to boninitic melts at shallow levels of a sub-arc mantle domain [1]. Our data yield age-corrected  $^{87}\text{Sr}/^{86}\text{Sr} = 0.7037\text{-}0.7045$  and  $\varepsilon_{\text{Nd}} = 0.3\text{-}7.5$ , interpreted as the result of mixing the depleted MORB mantle (DMM) and enriched mantle components (EM I and/or II). Along with strikingly unradiogenic initial Hf-isotope compositions ( $\varepsilon_{\text{Hf}} = -205\text{-}41$ ), this mixing may point to the contribution of an old and metasomatized mantle domain and/or subducted sediments.

Most internal and external Nd isochron ages (421-592 Ma), and second-stage Nd model ages (502-762 Ma) specify that the main magmatic episode occurred in Ordovician to Neoproterozoic times, in good agreement with a Re-depletion model age and Os isochron ages (588-801 Ma). We thus suggest that Cabo Ortegal pyroxenites recorded arc magmatism re-working a northern margin of Gondwana, or pre-Gondwanan continental blocks. In addition, based on Sr-isotope compositions of clinopyroxene and amphibole, we argue that previous ages reported at *ca* 390 Ma correspond to a metamorphic episode related to the syn-subduction exhumation of the Cabo Ortegal mantle, accompanied by extensive amphibolitization and metasomatic remobilization of Re, as recorded by the wide range of Os-isotope compositions ( $^{187}\text{Os}/^{188}\text{Os} = 0.16\text{-}1.44$ ) and  $^{187}\text{Re}/^{188}\text{Os}$ .

[1] Tilhac *et al.* (2016), *Journal of Petrology* **57**, 1921-1954.