## Multi-proxy evidence for global cooling during Oceanic Anoxic Event 2: the Plenus Cold Event

L.K. O'CONNOR<sup>1</sup>, S.A. ROBINSON<sup>1</sup>, S.J. BATENBURG<sup>1</sup>, H.C. JENKYNS<sup>1</sup>, B.D.A. NAAFS<sup>2</sup>, R.D. PANCOST<sup>2</sup>, & JWB RAE<sup>3</sup>

<sup>1</sup> Department of Earth Sciences, University of Oxford, South Parks Road, Oxford, UK

<sup>2</sup> Organic Chemsitry Unit, University of Bristol, Cantock's Close, Bristol, UK

<sup>3</sup> Department of Earth and Environmental Sciences, University of St. Andrews, Fife, UK

Oceanic Anoxic Event 2 (OAE2, ~94 Ma) was a period of extreme global warmth and represents one of the most extreme carbon cycle perturbations of the last 100 Myr [1]. The early stages of OAE2 are marked by a widespread episode of transient cooling and bottom water oxygenation known as the Plenus Cold Event. Originally recognised as a cooling event by the extreme southward migration of boreal fauna [2], a drop in temperature is supported by an excursion in  $\delta^{18}$ O and TEX<sub>86</sub> records from NW Europe and the North Atlantic [3,4,5]. The Plenus Cold Event is coincident with a marked decrease in *p*CO<sub>2</sub> [3] as well as a Nd-isotope excursion [6], indicative of changes in deep ocean circulation. However, these records are from a variety of separate localities and their temporal relation needs to be assessed critically.

Here we present preliminary results from an integrated study, in which we apply a range of geochemical proxy techniques on samples from the Plenus Marls cropping out in southern England. These diagenetically immature, organiclean nannofossil-rich chalks were deposited in an epicontinental pelagic setting and are ideal for palaeoclimatic reconstructions. Our high-resolution, multi-proxy reconstruction of the Plenus Cold Event will give greater insight into the timing and rate of change in sea-surface temperature, palaeo-pH, pCO<sub>2</sub>, as well as watermass behaviour across the event. The new data allow a detailed deconstruction of global feedback mechanisms operating during times of extreme global warmth.

[1] Jenkyns (2010) Geochem. Geophy. Geosy. 11. [2] Gale & Christensen (1996) B. Geol. Soc. Denmark. [3] Forster et al. (2007) Paleoceanography 22. [4] Jarvis et al. (2011) Paleoceanography 26. [5] Sinninghe Damsté et al. (2010) Earth Planet. Sci. Lett. 293, 97–103. [6] Zheng et al. (2013) Earth Planet. Sci. Lett. 293, 97–103.