

Magnesite stability at the reduced mantle conditions from the metal saturation depth to the lower mantle

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Magnesite is suggested to be a major oxidized carbon storage phase in the mantle due to its wide *P-T* range of stability [1]. However, the presence of magnesite in the Earth's interior depends highly on the redox conditions. Large part of the mantle is significantly reduced, with detectable amount of Fe⁰ dispersed in the silicate rocks [2]. Therefore, subducted carbonates would interact with Fe⁰. However, the mechanism of this interaction remains controversial.

We investigated MgCO₃-Fe⁰ interaction at 6–145 GPa and 800–2600 K. To determine influence of water on the reaction, additional experiments were conducted in the hydromagnesite-Fe⁰ system at 6 and 16 GPa. Redox reactions in both systems occur with the formation of (Mg,Fe)O, graphite/diamond, and Fe₃C/Fe₇C₃. The fundamental difference was connected with the reaction mechanisms. In the “dry” system solid state diffusion of components was suggested to be the major rate-limiting process with the reaction rate constant (*k*) in the range of 10⁻¹¹ -10⁻¹³ at 6 GPa and 1400 - 1600 K. In the water-bearing system, reactions were governed by solubility of components in the water fluids, and were controlled by the hydrogen fugacity. Water acts as a transporting agent for the reactants. Calculations of the kinetic parameters clearly detected the increase of *k* by two order of magnitude in comparison with the “dry” system.

The sluggish kinetics of the MgCO₃-Fe⁰ reaction established in our study suggests that carbonates could survive during subduction from metal saturation boundary near 250 km depth down to the transition zone and even to the core-mantle boundary. At the same time, in the water-containing systems magnesite would be completely reduced to carbide or diamond.

[1] Fiquet et al. (2002). *Am. Min.*, 87:1261-1265 [2] Frost., McCammon (2008). *Annu. Rev. Earth Planet. Sci. Lett.*, 36: 389-420.