## X-ray microtomography-based reconstruction of total CO<sub>2</sub> in olivine-hosted melt inclusions

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The global system of mid-ocean ridges represents a major locus of mantle CO2 outgassing [1]. Despite its importance to the global C budget, quantifying the CO2 flux from the mantle is hindered by magma degassing. The solubility of CO<sub>2</sub> in silicate melts is strongly pressure sensitive, leading to significant CO2 loss as magmas decompress and degas during ascent. Olivine-hosted melt inclusions (MIs) are an important source of information on the pre-eruptive CO2 contents of ocean floor lavas. The strength of the host crystal protects MIs from the decompression experienced by the entraining magma. However, a major uncertainty associated with using melt inclusions to determine the flux of CO2 from the oceanic upper mantle is the common occurrence of CO2-rich vapor bubbles (VB). In these instances, reconstructing total CO2 requires precise determination of the mass in the included glass and coexisting VB [2]. The concentration of CO<sub>2</sub> in the included glass can be measured using secondary ion mass spectrometry (SIMS), and the density of CO2 in the VB determined by Raman spectroscopy. To date, the volumes of included glass and coexisting VBs have relied on 2-D petrographic measurements mapped onto simplied 3-D geometries. We used x-ray microtomography to obtain precise determinations of these volumes. Results for olivinehosted MIs from pillow lavas of the late Pleistocene subglacial Miðfell eruption in south Iceland demonstrate that the 2-D petrographic approach systematically overestimates the MI volume and, thereby, underestimates total CO2. The concentration of CO2 measured in the glass portion of the MIs by SIMS ranges from 64 to 1160 µg/g. When present, vapor bubbles typically represent ~2.5 % of the volume of the inclusion, with CO<sub>2</sub> densities ranging from 0.06 to 0.25 g/cm<sup>3</sup>. Total CO<sub>2</sub> concentrations calculated for bubblebearing MI using 3-D tomographic volumes (1340-4550  $\mu$ g/g) are typically ~30 % higher than those determined using 2-D petrographic volumes (1010-3150 μg/g).

[1] Chavrit et al. (2014) *EPSL* **387**, 229-239. [2] Moore et al. (2015) *Am Min* **100**, 806-823.