

Decarbonation of subducting slabs: a petrological- thermomechanical modeling approach

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Subduction of heterogeneous lithologies carry a mixture of volatile components into the mantle. Volatiles can then be mobilized during metamorphic breakdown of hydrous and carbonate phases at shallow forearc depths followed by flux and carbonatite melting at subarc depths and beyond. While H₂O in subduction environments is studied extensively, processes resulting in CO₂ removal from the slab remain poorly constrained.

Mechanical removal of carbonated sediments via diapirism is hypothesized as an efficient mechanism to recycle CO₂ into the arc and mantle wedge. Therefore, we examine subduction decarbonation and sedimentary diapirism by employing a coupled petrological and thermomechanical numerical modeling approach. We modify the thermodynamic look-up tables [1] using *Perple_X* [2] to account for stable mineralogy and H₂O-CO₂ fluids within the following subducting lithologies: GLOSS average sediments (H₂O: 7.29 wt% & CO₂: 3.01 wt%), carbonated MORB (H₂O: 2.63 wt% & CO₂: 2.90 wt%), and carbonated peridotites (H₂O: 1.98 wt% and CO₂: 1.5 wt%). Our parametrizations include: 1) changing convergence velocity (1-6 cm year⁻¹; 1 cm year⁻¹ intervals) between a subducting oceanic plate and a continental plate and 2) thermal age of the subducting oceanic plate (20-80 Ma; 20 Ma intervals).

59 numerical models were run illustrating three geodynamic regimes that show extensive decarbonation. 1) Sedimentary diapirism acts as an efficient mechanism as it advects into the hot mantle wedge. CO₂ fluid is produced on the inner side of the advecting diapir. 2) If subduction rates are slow enough, mechanical coupling between the subducting oceanic plate and the continental plate results in a large mélange of fluid producing MORB and sediments. 3) During slab rollback and back-arc extension, interaction between hot asthenosphere and subducting mélange at shallow depths result in small windows (~12.5 Ma) of high integrated CO₂ fluxes (205 kg m⁻³ Ma⁻¹).

[1] Gerya *et al.* (2006) *PEPI* **156** 59-74. [2] Connolly JAD (2009) *Geochem. Geophys. Geosyst.* **10**:Q10014.