

Evanescent weathering surface evidenced by $^{234}\text{U}/^{238}\text{U}$ activity ratios?

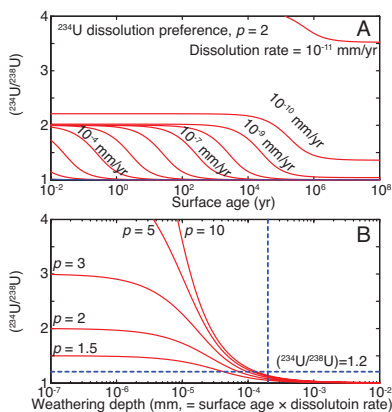
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Direct α -recoil and preferential dissolution of ^{234}U from the damaged lattice make ($^{234}\text{U}/^{238}\text{U}$) a potential rate-meter for mineral dissolution that is normalized to α -recoil scale of ~ 30 nm [1]. Here we present mathematical model and laboratorial experiments showing that the ($^{234}\text{U}/^{238}\text{U}$) of weathering solution is functions of surface age, dissolution rate, and dissolution preference (p) of the ^{234}U sites (Fig. 1A). For a given rock type with fixed p value, the ($^{234}\text{U}/^{238}\text{U}$) is approximately a monotonic function of dissolution depth if dissolution rate is higher than 10^{-9} mm/yr (Fig. 1B). The ($^{234}\text{U}/^{238}\text{U}$) value of basaltic catchments [2-7], which is mostly higher than 1.2, indicating dissolution depth of less than $0.2 \mu\text{m}$ (Fig. 1B), and thus implying an evanescent weathering surface of about twenty years given a weathering rate of $\sim 10^{-2}$ mm/ky [8]. Accumulation of leached layers and secondary precipitates may inhibit deep dissolution. Interestingly, the depth of mineral dissolution, as proxied by riverine ($^{234}\text{U}/^{238}\text{U}$), correlates well with weathering temperature.



[1] Vigier, et al., 2001, *EPSL*, **193**, 549; [2] Riotte et al. 2003, *Chem. Geol.* **202**, 365; [3] Vigier et al., 2005, *Chem. Geol.* **219**, 69; [4] Vigier et al., 2006, *EPSL*. **249**, 258; [5] Pogge von Strandmann et al., 2006, *EPSL*, **251**, 134; [6] Pogge von Strandmann et al., 2010, *Chem. Geol.*, **270**, 227; [7] Bagard et al., 2011, *GCA*, **75**, 3335; [8] Navarre-Sitchler, et al., 2007, *EPSL*, **261**, 321.