High-temperature high-pressure carbon isotope fractionation between COH-fluids and elemental carbon

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Carbon isotope fractionation in the system carbon-oxygenhydrogen-fluid (COH-fluid) - graphite/diamond at deepcrustal and mantle conditions is currently best described by theoretical calculations. These predict a fractionation of -1‰ to 5‰ between CH4 or CO2 and graphite/diamond at temperatures between 800 to 1400°C [1-3]. Experimental work verifying these calculations is scarce. Our goal is the determination of equilibrium carbon isotope fractionation factors between elemental carbon (C^0) and COH-fluid at 800 to 1600°C, 2 to 60 kbar and varying oxygen fugacities at least between QFM and IW-3. The O:H ratio of the starting material defines the fluid speciation, i.e. CH4, CO, and CO2 fractions and the fO2-value of the experiments [3] (Fig.1) and is therefore crucial for C-isotope fractionation [4]. We have hence designed a capsule-piercing-cell-assembly that allows contamination-free gas extraction from the experimental capsules and gas-storage in vials permitting multiple isotope and speciation analysis of the recovered fluid. The first experiments run in a fast-quench externally-heated pressure vessel at 2 kbar and 800°C produced sufficient volumes of gas and C⁰ for multiple isotope analysis.

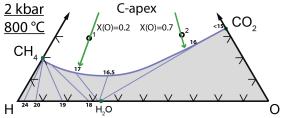


Figure 1: COH-ternary diagram. Blue C⁰-saturation curve separates the C⁰-COH-fluid field (white) from the COH-fluid field (gray). fO_2 -isopleths are labeled in $-log(fO_2)$. COH-compounds 1 and 2 decompose to a CH₄ and a CO₂-dominated fluid that would be in equilibrium with C⁰.

[1] Bottinga (1969) Geochim. Cosmochim. Acta 33, 49-64. [2]
Bottinga (1969) Earth Planet. Sci. Lett. 5, 301-307 [3]
Connolly & Cesare (1993) J. Metamorph. Geol. 11, 379-388.
[4] Deines (1980) Geochim. Cosmochim. Acta 44, 943-961.