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Magnetite-apatite ore bodies of the Kiruna type occur worldwide and are associated with volcanic rocks, which suggests that they represent highly evolved, late-stage, igneous melts/bodies, associated with volcanism, either as surface lavas or as shallow intrusions, with a later metasomatic component. However opinion varies with regard to this origin with a sedimentary origin suggested for some ore bodies (e.g. see discussion in Naslund et al., 2002, Porter, T.M., ed., Hydrothermal Iron Oxide Copper-Gold & Related Deposits: A Global Perspective, volume 2, PGC Publishing, Adelaide, pp. 207–226). Notable examples of Kiruna type magnetite-fluorapatite ore bodies include the Kiirunavaara ore body, northern Sweden (Harlov et al., 2002a, Chem. Geol., 191, 47-72); the Mineville ore body, Adirondacks, New York, USA (McKeown and Klemc, 1956, U.S. Geol Sur Bull (1956), pp. 9-23); the Pea Ridge ore body, Arkansas, USA (Kerr, 1998, MS Thesis, University of Windsor, 113 pp); a series of ore bodies from the Bafq Region, central Iran (Torab and Lehmann, 2007, Mineral. Mag., 71, 347–363); and a series of ore bodies associated with volcanic activity in the Plio-Pleistocene Chilean High Andes and the Cretaceous Coastal Andean Cordillera (Naslund et al., 2002).

In each of these ore bodies (except the last example) monazite +/- xenotime inclusions are commonly found in areas of the fluorapatite, which have experienced obvious fluidinduced alteration in the form of (Y+REE) + Na + Si + Cl depletion (e.g. Kiirunavaara, northern Sweden; Harlov et al., 2002a) with the fluorapatite serving as the source of P + (Y+REE). Formation of monazite and xenotime inclusions in fluorapatite has experimentally been demonstrated to result from fluid aided coupled-dissolution-reprecipitation processes (Harlov et al., 2005, Contrib. Mineral. Petrol. 150, 268-286). Fluid sources could range from 800-900 °C, residual acidic (HCl, H₂SO₄) grain boundary fluids remaining after the last stages of ore body crystallization to later stage, cooler (H2O-CO₂-(Na,K)Cl) fluids originating in the surrounding country rock or fluids associated with events such as regional albitization or actinolization. Re-integrating $(Y+REE)PO_4$ inclusions back into the fluorapatite indicate original $(Y+REE)_2O_3$ contents ranging from 3 – 5 wt% suggesting that in addition to being mined for their iron ore, the fluorapatite in these ores could be mined for (Y+REE) as well.