## Putting all the letters in *P-T-A-X-d-t*

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Ages of mineral chronometers can vary as a function of:

Pressure P. Diffusivity D decreases when P increases. Small effect, compared to the other letters.

Temperature T. Everything goes faster when T increases. Just how important this is will be seen below.

Water activity A. It enables dissolution-reprecipitation, without which no medium-T metamorphic paragenesis can form [1]. It overruns every other process and has a very weak temperature dependence [2].

Molar fraction X. It determines the stability fields in pseudosections. Destabilized mineral chronometers are reset, stable ones mostly not [3].

Deformation d. By promoting recrystallization it resets mineral ages at low T while samples a few m away preserved old ages at high T [4,5].

Age *t*. Intra-mineral age variations recognized by spatial resolution and/or stepwise release ("SR/SR") help us understand the respective roles of P,T,A,X,d. [6]

Diffusive losses can occur in some cases [3,7] and are diagnosed by characteristic intra-grain bell-shaped profiles [8,9]. Irregular textures require high A independently of T. As chronometers record the event that reset them most efficiently, patchy minerals are a poor thermochronometer but a good hygrochronometer. The special role of petrologic equilibrium [10] is easily summarized: equilibrium = maybe diffusion, disequilibrium = certainly retrogression & recrystallization. Partly retrogressed phases (orthoclase/adularia, phengite/paragonite/muscovite, biotite/chlorite, allanite/monazite, etc) require not only age acquisition by SR/SR but above all the diagnose of heterochemical replacement [11, 6].

Retrogression reactions depend on PTAXd; the dependence of t on T is subordinate and never unique.

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