Mixing of solid precursors during formation of FeO-poor chondrules in CR3 chondrites

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In CR3 chondrites, FeO-poor chondrules measured by SIMS [1] have $\delta^{\rm 18,17}O$ values on the slope ${\sim}1$ primitive chondrule mineral (PCM) line [2]. They are ¹⁶O-poor relative to the bulk Solar System estimate [3], but ¹⁶O-rich relative to O associated with H₂O [4] and organics [5] in CR chondrites. Δ^{17} O values systematically increase from -6 to ~-1 ‰ as chondrule Mg# decreases from 99.2 to ~96. The Mg# range implies fO_2 varied by 1.1 log units in the chondrule-forming environment (IW -3.6 to IW -2.5, respectively; e.g. [6]), suggesting addition of oxidizing ¹⁶O-poor H₂O to the highest Mg# chondrule precursors as a mechanism to form lower Mg# chondrules [7]. Using existing equilibrium condensation models and an O-isotope mass balance we examine if the CR3 chondrule Mg#- Δ^{17} O trend can be explained by precursors variably mixed in Solar gas and a CI chondrite composition split into silicate, organic, and H₂O components. In the model, ratios of atomic H, O, and C, in assemblages, dependent on the degree of enrichment in dust, H_2O_1 , and organics, determine fO_2 during chondrule-formation [e.g. 6]. We assigned O-isotope ratios of precursor components, based on measurements/estimates [1,3,5], allowing for quantifying model parameters that match the observed CR3 chondrule trend. Our model predicts FeO-poor CR3 chondrules formed at dust enrichments of 100-200×, from dusts with 0 to 0.8 times the amount of H₂O in CI chondrites

Tenner et al (2012) 43rd LPSC #2127. [2] Ushikubo et al (2012) GCA 90, 242-264. [3] McKeegan et al (2011) Science 332, 1528-1532. [4] Clatyon & Mayeda (1999) GCA 63, 2089-2104. [5] Hashizume et al (2011) Nature Geosci. 4, 165-168. [6] Ebel & Grossman (2000) GCA 64, 339-366. [7] Connolly & Huss (2010) GCA 74, 2473-2483.