

## Mantle sources of Archean gabbroanorthosite of the Fennoscandian Shield: isotopic study of U-Pb in zircon and isotope-geochemistry data

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Archean gabbroanorthosite magmatism is located in the north-eastern part of the Fennoscandian Shield and is presented by Neoproterozoic (2.7-2.6 Ga) [1] and Mesoproterozoic intrusions (2.9 Ga).

The Mesoproterozoic gabbroanorthosite intrusions are similar to the anorthosites of the Fiskeneset Complex (Greenland), the Karnataka craton (India) and the Vermillion Lake Complex (Canada). U-Pb zircon dating established Mesoproterozoic ages of  $2925 \pm 7$  and  $2935 \pm 8$  Ma for the gabbroanorthosites of the Patchemvarek and Severny intrusions, respectively. Zircon has typical igneous REE patterns and oscillatory zoning. The normalized pattern is characterized by a steeply-rising slope from the LREE to the HREE with a positive Ce-anomaly and negative Eu-anomaly.

It was shown that the gabbroanorthosite have fairly low REE contents ( $Ce_n = 2.2-4.2$ ,  $Yb_n = 1.6-2.6$ ) and distinct positive Eu anomaly. Comagmatic ultrabasic differentiates have practically unfractionated REE pattern, low total REE contents ( $Ce_n = 1.2$ ,  $Yb_n = 1.1$ ,  $La/Yb_n = 1.32$ ), and have not Eu anomaly. The Mesoproterozoic gabbroanorthosites are characterized by positive  $\epsilon_{Nd} = +2.77 - +1.66$  and Neoproterozoic gabbroanorthosites -  $\epsilon_{Nd} = +1.78 - +0.26$ . The Sm-Nd isotope data suggest the existence of several mantle sources in the NE Fennoscandian Shield, which produced melts for different-age gabbroanorthosite.

The Mesoproterozoic gabbroanorthosites were presumably derived from MORB-type basalts of oceanic settings, while the Neoproterozoic gabbroanorthosites were generated from subalkaline magma formed within plate anorogenic setting.

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[1] Bayanova T.B. Vozrast reperykh geologicheskikh kompleksov Kol'skogo regiona i dlitel'nost' protsessov magmatizma (Age of Reference Geological Complexes of the Kola Region and Duration of Magmatic Processes), St. Petersburg: Nauka, 2004.