Oxygen isotope systematics of chondrules among chondrite groups

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Oxygen isotope ratios of chondrules record those of the dust-enriched protoplanetary disk in the early Solar System, from which their parent asteroids accreted [1-2]. High precision SIMS oxygen isotope analyses of chondrules in the least equilibrated chondrites have been conducted on >300 chondrules from various chondrite groups; LL3, H3, Acfer 094 (ungrouped C), CO3, CV3, CR3, CH3, and R3 [1-9].

Oxygen isotope ratios of chondrules in carbonaceous chondrites [2,4-8] are characterized by enrichment of ¹⁶O relative to the terrestrial mantle. Their δ^{18} O and δ^{17} O values plot along the slope ~1 primitive chondrule mineral (PCM) line, defined in [2]. They show a wide range of Δ^{17} O, mainly from -6‰ to +1.5‰, and correlate with Mg#. The slope ~1 chondrule trend could be the result of mixing of ¹⁶O-rich anhydrous dust and ¹⁶O-poor ice in precursor solids within the dust-enriched disk [10], which could be related to extreme mass independent isotope fractionation (MIF) in the earliest solar nebula, as indicated from [11-12].

In contrast, ordinary and R chondrite chondrules show mass dependent isotope fractionation above the terrestrial fractionation line ($\Delta^{17}O \sim +1\%$), toward lower $\delta^{18}O$ values relative to the PCM line [1,3,9]. FeO-poor olivine-rich chondrules (IA) have the lowest $\delta^{18}O$ values [1]. A narrow range of $\Delta^{17}O$ values could indicate isotope reservoirs were well-mixed prior to the chondrule forming event(s). [1] argued that a significant isotope fractionation ($\geq 3 \%$) could be produced between nebula gas and olivine without dust enrichment at ~1200°C. MIF components in inner asteroidal regions could have been homogenized by transient heating events early during disk evolution, before dust accreted and densified.

[1] Kita N. T. et al (2010) GCA 74, 6610-6635 [2] Ushikubo T. et al (2012) GCA 90, 242-264 [3] Kita N. T. et al (2008) MaPS
43, A77 [4] Rudraswami N. G. et al (2011) GCA 75, 7596-7611 [5] Tenner T. J. et al (2013) GCA 102, 226-245 [6] Tenner et al T. J. (2012) 43rd LPSC #2127 [7] Nakashima D. et al (2010) MaPS 45, A148 [8] Nakashima D. et al (2011) MaPS 46, 857-874 [9] Kita N. T. et al (2013) MaPS 78, 5149 [10] Tenner T. J. et al (2014), this meeting. [11] Sakamoto N. et al (2007) Science 317, 231-233 [12] McKeegan K. D. et al (2011) Science 332, 1528-1532