

## Paleoarchean felsic magmatism: A melt inclusion study of 3.45 Ga volcanic rocks from the Barberton Greenstone Belt

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Archean felsic magmatism spatially and temporally associated with the Barberton Greenstone Belt (BGB) of Southern Africa can be broadly subdivided in an early ( $\geq 3.2$  Ga), Na-rich series that formed the tonalite-trondhjemite-granodiorite (TTG) series, and a generally later ( $\sim 3.1$  Ga), K-rich series akin to modern granites [1].

Felsic volcanic rocks in the BGB are less abundant than their intrusive counterparts, and are mostly strongly deformed. The 3.45 Ga old Buck Ridge Volcanic Complex (BRVC; [2]) contains the least deformed felsic volcanic rocks in the BGB. However, widespread alteration (silicification, K-metasomatism) of these rocks involved strong mobilisation of major and some trace elements [3].

Quartz-hosted melt inclusions have been used to obtain information on the unaltered melt composition. Whole-rock compositions of the least altered samples bear strong resemblance with coeval TTG intrusions, and can be modeled as a mixture of modal phenocrysts (Na-plagioclase, quartz, Fe-Ti oxide, apatite) and melt as indicated by melt inclusion analyses, thus suggesting a cogenetic relationship between the BRVC and TTG intrusions. This is also confirmed by whole-rock concentrations of fluid-immobile trace elements. Further, melt inclusions have moderate Cl contents ( $\leq 0.7$  wt.%), low F and low S ( $\leq 0.12$  and  $< 0.02$  wt.%, respectively). The moderate Cl and low F/Cl suggest a sea water contribution to the melting protolith [4]. These characteristics are interpreted as due to melting of source rocks that have undergone sea floor metamorphism, although in an Archean context this does not have immediate tectonic implications.

[1] Moyen, 2011. *Lithos* **123**, 31-36. [2] de Vries *et al*, 2006. *Prec. Res.* **149**, 77-98. [3] Hofmann and Harris, 2008. *Chem. Geol.* **257**, 221-239 [4] Pyle and Mather, 2009. *Chem. Geol.* **263**, 110-121.

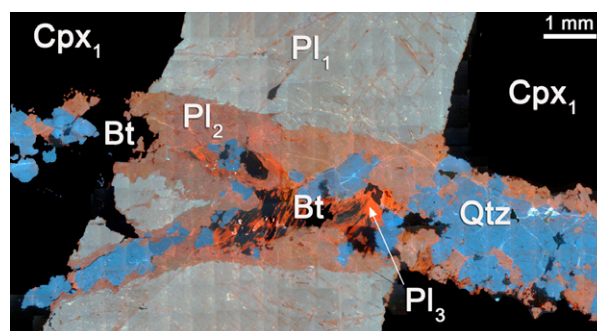
## Magmatic and hydrothermal history of felsic rocks within oceanic core complex (MAR, 13°31'-13°35' N)

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Felsic plutonics and dykes – oceanic plagiogranites (OPG) – spatially associate with inactive hydrothermal fields and are regarded as the products of gabbro/dolerite hydrous melting within the oceanic core complex at 13°31'-13°35' N, Mid-Atlantic Ridge. We used cathodoluminescence (CL) and SIMS analytics to reveal mineralogical and geochemical features, and distinguish main stages of the OPG formation.

The OPG lithology includes: metadolerite - amphibolite veined by hornblende plagiogranite; coarse-grained gabbro with biotite-bearing plagiogranite veins; massive hornblende and amphibole-free OPG. The CL image (Fig. 1) demonstrates relationships between mineral assemblages, corresponding to three successive stages: mafic igneous ( $Pl_1$ ,  $Cpx_1$ ), felsitic igneous ( $Qtz$ ,  $Bt$ ,  $Pl_2$ ,  $Cpx_2$ ,  $Zrn$ ) and hydrothermal ( $Qtz$ ,  $Pl_3$ ,  $Chl \pm Cpx_3$ ).



**Figure 1:** The OPG-vein in gabbro (CL image).

The trace-element data allow distinguishing: (1) mafic (gabbroic) fractionation; (2) felsic (OPG) fractionation; and (3) hydrothermal processing.  $Pl_1$  (An65-60) and  $Cpx_1$  show core-to-rim REE enrichment being appreciably depleted as compared to  $Pl_2$  and  $Cpx_2$  respectively. Felsic stage associates with subsequent REE enrichment of  $Pl_2$  (An47-27) and  $Cpx_2$ . The hydrothermal stage exhibits  $Pl_3$  (An17-5) and  $Cpx_3$  that are significantly depleted in REE.