

Atmospheric trace gases and isotopologues using mid-IR laser direct absorption spectroscopy

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Recent advances in mid-infrared laser technology have greatly facilitated field measurements of atmospheric trace gases to identify and quantify their sources and sinks. Newly available non-cryogenic mid-infrared lasers and detectors, long path length, small volume sampling cells, and advances in direct absorption techniques with have led to smaller, lighter, and more robust instrumentation for measurements from mobile and aircraft platforms, field sites at remote locations, and laboratories for clumped isotopologues. The reliability and reproducibility of the mid-infrared lasers has led to long term monitoring and turn-key operation. Trace gas detection in ambient air at the low part-per-trillion levels are now feasible. Fractional precisions of less than 1 part in 10,000 allow for isotopologue ratio measurements of carbon dioxide, methane, and nitrous oxide at atmospheric mixing ratio levels. Applications to measurements of greenhouse gas emissions of methane and nitrous oxide, high precision measurements of carbonyl sulphide (OCS), measurements of ethane (C₂H₆) to determine sources of methane, and isotopologue measurements of CO₂, CH₄, N₂O and H₂O, will be presented.

TRACE GASES ambient mixing ratios	Frequency cm ⁻¹	1 s std dev [ppt] 210 m 76 m	
CH ₄	1275	100	300
¹³ CH ₄ /CH ₄	1294	1.5‰	
OCS	2050	2	5
CO	2199	40	
N ₂ O	2199	20	
¹⁵ N ¹⁴ NO, ¹⁴ N ¹⁵ NO	2188	3‰	
¹³ CO ₂ , C ¹⁸ O ¹⁶ O	2311	0.1‰	
C ₂ H ₆	2996	20	
¹³ CH ₄ /CH ₄	3057	0.5‰	
CH ₃ D/CH ₄	3060	20‰	

Impacts and LIPs: ¹⁸⁷Os/¹⁸⁸Os signatures across the K-Pg boundary

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Both impacts and large igneous provinces (LIPs) can introduce unradiogenic Os into the global ocean producing excursions to lower ¹⁸⁷Os/¹⁸⁸Os in pelagic sediments. The ¹⁸⁷Os/¹⁸⁸Os excursion caused by the late Eocene Popigai impact is consistent with abrupt addition of soluble meteoritic Os to the ocean [1]. In contrast, previously reported ¹⁸⁷Os/¹⁸⁸Os excursions related to LIPs show greater variability and longer duration than the Popigai Os excursion. It is well established that the Chicxulub impact event is coincident with the K-Pg mass extinction and that eruption of the Deccan Traps began before, and continued after, the extinction event. Comparison of the Popigai excursion to the Cretaceous-Paleogene (K-Pg) boundary ¹⁸⁷Os/¹⁸⁸Os record provides a means of assessing whether or not eruption of the Deccan liberated significant amounts of mantle-derived Os to the ocean-atmosphere system coincident with the K-Pg mass extinction. In the Equatorial Pacific, South Atlantic and Atlantic-sector of the Southern Ocean the recovery of the marine ¹⁸⁷Os/¹⁸⁸Os record to higher pre-extinction levels [2] closely resembles the shape and duration of the Os recovery from the Popigai impact event. This similarity does not support claims [3] that the main phase of the Deccan volcanism is closely linked to the K-Pg mass extinction. However, an Indian Ocean site (ODP 738C) on the Kerguelen Plateau displays ¹⁸⁷Os/¹⁸⁸Os ratios consistently below those from other ocean basins during the Paleogene. In close proximity to the K-Pg boundary in ODP 738C, the Os-Ir-¹⁸⁷Os/¹⁸⁸Os signature is unequivocally meteoritic. Upsection, low ¹⁸⁷Os/¹⁸⁸Os ratios as far as 4.5m above the boundary (≈ 1.2 M.Y.) are unlikely to be related to the Chicxulub impact event. These low ¹⁸⁷Os/¹⁸⁸Os ratios in the Indian Ocean may be fingerprint of Paleogene LIP activity associated with either the Deccan or Kerguelen, but other interpretations are possible. The marine ¹⁸⁷Os/¹⁸⁸Os record allows large impact events to be differentiated from the emplacement of LIPs, and may preserve regional gradients in seawater ¹⁸⁷Os/¹⁸⁸Os controlled by the proximity to active volcanism.

[1] Paquay *et al.* (2008) *Science* 320, 214-218. [2] Ravizza and Vonderhaar (2012) *Paleoceanography* 27, 10.1029/2012PA002301. [3] Keller *et al.* (2009) *EPSL* 282, 10-23.