

## Petrogenesis of the Syenite Granites in Kuluketage Block: Constraints from Petro-Geochemistry, Zircon U-Pb Dating and Hf Isotope

Q. YUAN, X.B. LÜ\*, X.F. CAO, X.D. WANG  
AND E.L. YANG

Faculty of Earth Resources, China University of Geosciences,  
Wuhan 430074, China  
(\*correspondence: lvxb\_01@163.com)

Kuluketage block is the best area for Precambrian geology in Xinjiang, NW China. In this paper, we studied the petrology, geochemistry, zircon LA-ICPMS U-Pb chronology and zircon Hf isotope research of Daxigou syenite granite (DSG) which is located in Kuluketage. Zircons from the intrusions display oscillatory zoning and high Th/U ratios (0.16-2.11), implying their magmatic origin. Zircon U-Pb dating results of DSG indicate that they formed in Paleoproterozoic with the weighted  $^{207}\text{Pb}/^{206}\text{Pb}$  average age of  $1767 \pm 46 \text{ Ma}$ , which is coincidentally identical with its associated diorite age within the error range. Studies of petrogeochemistry suggest that DSG belong to medium-sodium and rich-potassium peraluminous calc-alkaline type, rich in Pb, La, Th and LILE, significant poor in HFSE (Gd, Nd, Ta). The chondrite-normalised REE pattern is slightly to the right form. The average  $\Sigma \text{ REE}$  is  $56.57 \times 10^{-6}$ ; HREE show moderate fractionation (LREE/HREE averaged 10.28, (La/Yb)<sub>N</sub> average of 13.04) N; average (La/Sm)<sub>N</sub> of syenite is 4.96, average (Gd/Yb)<sub>N</sub> is 1.67 and the  $\delta \text{Eu}$ ,  $\delta \text{Ce}$  are not obvious. Their initial Hf isotope ratios and Hf two-stage model ages range from -7.74 to -4.02 and 2.74 Ga to 2.67 Ga, respectively. Taken together, it is suggested that DSG is the typical S-type granite and its primary magma could be mainly derived from partial melting of the Neoproterozoic crust and mainly formed in the syn-collision arc environment, which recorded the tectonic-magma activities response of Tarim refers to the amalgamation of the supercontinent Columbia.

## Copper-mediated oxidation of hydroquinone under conditions typical of natural saline waters

XIU YUAN<sup>1</sup>, A. NINH PHAM<sup>2</sup>, CHRISTOPHER J. MILLER<sup>3</sup>  
AND T. DAVID WAITE<sup>4</sup>

School of Civil and Environmental Engineering, the  
University of New South Wales, Sydney, NSW 2052,  
Australia

<sup>1</sup> (xiu.yuan@unsw.edu.au.) presenting author;

<sup>2</sup> (anninh.pham@unsw.edu.au.); <sup>3</sup> (c.miller@unsw.edu.au.);

<sup>4</sup> (d.waite@unsw.edu.au.)

Copper (Cu) is an essential transition metal that is involved in a variety of photoreactions and physiological processes. The redox chemistry between cuprous copper (Cu(I)) and cupric copper (Cu(II)) in the upper water column plays a significant role in its speciation, transport and bioavailability [1]. As electron transfer mediators, quinone moieties in natural organic matter (NOM) play important roles in essential biogeochemical processes [2-3]. In this study, detailed kinetic model has been developed to describe the oxidation of 1,4-hydroquinone (H<sub>2</sub>Q) by Cu(II) in the presence O<sub>2</sub> in 0.7 M NaCl solution at pH range 6.5 – 8.0. Cu(II) catalyzed the overall oxidation of H<sub>2</sub>Q in a strongly pH dependent manner with concomitant formation of benzoquinone and H<sub>2</sub>O<sub>2</sub>.

The kinetic model indicated the mono-anion HQ<sup>-</sup> is the active species to reduce Cu(II) with an intrinsic rate constant of  $5.0 \times 10^7 \text{ M}^{-1} \cdot \text{s}^{-1}$ . While the semiquinone radical (SQ<sup>-</sup>) generated from one-electron oxidation of hydroquinone or one-electron reduction of benzoquinone acts as a chain propagating species, the presence of O<sub>2</sub> in the system is also essential in terms of supplying Cu(II) by continuously oxidizing Cu(I) and rapidly removing SQ<sup>-</sup> to generate O<sub>2</sub><sup>-</sup>. The development of the kinetic model should assist in understanding and predicting the factors controlling copper transformation, speciation, bioavailability and toxicity in aquatic systems.

[1]. Moffett, J. W.; Zika, R. G. (1988) *Geochim. Cosmochim. Acta* 52, 1849-1857. [2]. Uchimiya, M.; Stone, A. T. (2010). *Aquat Geochem* 16, 173-188. [3]. Uchimiya, M.; Stone, A. T. (2009), *Chemosphere* 77, 451-458.