

The HP stability of Bloedite ($\text{Na}_2\text{Mg}(\text{SO}_4)_2 \cdot 4\text{H}_2\text{O}$): a contribution to the knowledge of asteroids and icy satellites

PAOLA COMODI^{1*}, SABRINA NAZZARENI¹, TONCI BALIC-ZUNIC², AZZURRA ZUCCHINI¹, MICHAEL HANFLAND³

¹ Dipartimento Scienze della Terra, Università di Perugia, Italy;
comodip@unipg.it (* presenting author)

² Natural History Museum of Denmark, University of Copenhagen,
Denmark, TonciB@snm.ku.dk

³ ESRF, Grenoble, France, hanfland@esrf.fr

Bloedite, $\text{Na}_2\text{Mg}(\text{SO}_4)_2 \cdot 4\text{H}_2\text{O}$ is a common mineral in evaporitic marine sediments. Being one of the phases in the system H_2O - MgSO_4 - NaSO_4 it has a large planetological interest. In fact volatiles of icy satellites are considered to contain brines with sulphates, as well as carbonates and chlorides. To understand their internal structure and to evaluate the presence of a deep internal ocean, it is important to know the phase relation between water and brine mixtures [1]. Bloedite crystal structure consists of parallel layers of $\text{MgO}_2(\text{H}_2\text{O})_4$ and $\text{NaO}_4(\text{H}_2\text{O})_2$ octahedra, interconnected through SO_4 tetrahedra and hydrogen bonds [2]. This paper intends to investigate the high-pressure behavior of bloedite to determine the equation of state, the density evolution, as well as the dehydration conditions. Data were collected at ID-09 beamline, at ESRF (Grenoble, France). We used a diamond anvil cell with 300 micron diamond culet, Neon gas as pressure transmitting medium and ruby chip as pressure calibrant. The very large image-plate detector, Mar555, allowed to collect data of high quality from single crystals. Lattice parameters and reflections intensities were obtained by CrysAlis software. SHELXL software was used to refine the structure at different pressures up to 12 GPa. Equation of state (EoS) of bloedite was determined with data collected up to 12 GPa and EOS-Fit program [2]. A second order Birch-Murnaghan EoS fit yields $V_0 = 495.6(9) \text{ \AA}^3$ and $K_0 = 40.5(7) \text{ GPa}$, whereas a third order Birch-Murnaghan EoS fit yields $V_0 = 497.6(4) \text{ \AA}^3$, $K_0 = 34(1) \text{ GPa}$ and $K' = 5.8(4) \text{ GPa}^{-1}$. The lattice parameters compressibilities are $\beta_a = -0.0074(4) \text{ GPa}^{-1}$; $\beta_b = -0.0069(4) \text{ GPa}^{-1}$; $\beta_c = -0.0054(2) \text{ GPa}^{-1}$ with an anisotropic ratio of 1: 1.07: 1.37. The structure is more incompressible along the direction perpendicular to open sheet, than the other directions. The SO_4 tetrahedra are incompressible, K_{MgO_6} octahedra is 90 GPa and K_{NaO_6} octahedra is 42 GPa. No phase transitions were observed in the bloedite structure up to 12 GPa, who remained stable at very HP. Hydrogen bonds evolution were followed through the $\text{O}_{\text{donor}}\text{-O}_{\text{acceptor}}$ distance, whose configuration explained the large compressibility of a and b parameters with respect to c lattice parameter. The lack of strong structural rearrangement, essential to compensate the dehydration process, suggested that water remained in the structure of bloedite in the investigated P range.

[1] Nokamura, Othani (2011) *Icarus*, 211, 648-656, [2] Hawthorne (1985) *Canadian Mineralogist*, 23, 669-674, [3] Angel (2000) *Reviews in Mineralogy and Geochemistry* 41.

Syenitic Provinces in the São Francisco Craton, Brazil

HERBET CONCEIÇÃO^{1*}, MARIA LOURDES SILVA ROSA¹,
DÉBORA CORREIA RIOS²

¹Sergipe Federal University, Núcleo de Geologia, Aracaju, Brazil,
herbet@ufs.br

²Bahia Federal University, Instituto de Geociências, Salvador, Brazil,
dcorios@ufba.br

The Paleoproterozoic in the Bahia State, Brazil, is characterized by the generation and emplacement of various syenite bodies. The mainly occur in two mobile belts that represent part of the basement of the São Francisco Craton. These mobile belts are: (i) the Salvador-Curaçá mobile belt (SCMB), located in the eastern part of the state, and (ii) the Urandi-Paratinga mobile belt (UPMB), in the western part. Both are characterized by the emplacement of alkali-syenitic, potassic to ultrapotassic rocks, during the late stages of the stabilization of these mobile belts, which occurred between 2.1 and 2.0 Ga. However, in the east part the syenite is placed as tabular bodies and shows gneissic texture, while in the west part they are related to pull-part system and develop only magmatic textures. The UPMB syenite occurs as wide single batholith (6,000 km^2), the Guanambi Batholith (GB), dated $2.05 \pm 0.02 \text{ Ga}$ (U-Pb age). In the GB two main domains were recognized; (i) multiple intrusions (about 92% of the batholith), and (ii) late intrusions. Both have composition from syenite to mafic syenite towards monzonite. The U-Pb data show that the difference between the emplacement of late intrusions was less than 5 Ma. The Eastern Bahia State, the SCMB syenite consists mainly of four massifs, discontinuously disposed. From north to south occur the following bodies: Itiúba (1800 km^2 , $2.00 \pm 0.03 \text{ Ga}$, Pb/Pb age), Santanópolis (180 km^2 , $2.10 \pm 0.04 \text{ Ga}$, U-Pb age), São Felix (32 km^2 , $2.09 \pm 0.01 \text{ Ga}$, Pb-Pb age) and Anuri (70 km^2 , $2.10 \pm 0.020 \text{ Ga}$ U-Pb age). All intrude both gneiss-migmatitic and granulitic terrains. These syenites, despite the mobile belt in which they occur, show the same petrographic characteristics. They are leucocratic, porphyritic, mainly composed of perthitic alkali-feldspar, diopside, hornblende, phlogopite and biotite, ilmenite and magnetite. In the syenites was found lamprophyric dykes (minette). Geochemical data indicate that all these syenites are SiO_2 -saturated to oversaturated, alkalic to sub-alkalic and metaluminous. The $\text{K}_2\text{O}/\text{Na}_2\text{O}$ ratios are always greater than unity, being higher in mafic terms. In some cases, these mafic terms can be classified as ultrapotassic. Ba (up to 8,000 ppm), Sr (up to 6,000 ppm) and Rb (up to 940 ppm) are strongly enriched in these rocks. Cr (from 50 to 700 ppm), Ni (from 80 to 270) and Mg# (from 0.20 to 0.77) are relatively high for common syenitic rocks. Chondrite-normalized REE patterns show strongly fractionated LREE with a small negative Eu anomaly, $\epsilon_{\text{Nd}(t)}$ range from -10 to 0 (SCMB), and -11 to -7 (UPMB) and Sr_i values are around 0.705. From those data, we can conclude that the Paleoproterozoic was a propitious time for the generation of syenitic rocks in these two mobile belts of the Bahia State, independently from the tectonic regime. The syenitic result probably from fractionated crystallization of a lamprophyric magma, the later generated by melting of Paleoproterozoic enriched mantle. *Acknowledgments: This work was supported by CNPq, CBPM and FAPITEC.*