## High-precision <sup>26</sup>Al-<sup>26</sup>Mg dating solid and planetesimal formation in the young Solar System

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The short-lived  $^{26}$ Al-to- $^{26}$ Mg isotopic system ( $t_{1/2} = 0.73$ Myr) is potentially the most powerful chronometer to study early Solar System timescales. This reflects the short half-life of <sup>26</sup>Al and large Al/Mg fractionations induced by nebula and planetary processes, which coupled with the precisely determined initial <sup>26</sup>Al abundance of CAIs, makes it the only isotopic system that allows relative age determinations on meteorites to be made with a resolution of ±0.1 Myr. However, full utilization of the <sup>26</sup>Al-<sup>26</sup>Mg dating system requires the measurement of  $\delta^{26}$ Mg\* excesses or deficits with an accuracy and precision of ±0.003 to 0.004% (2 se). We have developed new methods for measuring the Mg isotope ratios of meteorite material by pseudo-high resolution MC-ICP-MS. Tests have been conducted to ensure that  $\delta^{26}$ Mg\* can be determined with an accuracy and precision of  $\leq \pm 0.004\%$ . Improved chemical separation techniques have been developed to minimise the contaminant elements (e.g. Mn, Ni) in purified Mg, which are not separated by conventional nitricacid-based cation exchange techniques.

Data obtained thus far shows: (1) Unaltered bulk CAIs from CV chondrites have resolvable differences in initial <sup>26</sup>Al/<sup>27</sup>Al from 4.7 to 5.2 x 10<sup>-5</sup> (corresponding to a time interval of 89±30 kyr) with no CAIs showing evidence for "supercanonical" <sup>26</sup>Al. (2) Most basaltic meteorites from the eucrite (3.1 Myr), mesosiderite (2.8 Myr) and angrite (4.8 Myr) parent bodies have elevated  $\delta^{26}$ Mg\* that dates planetary melting to 2.8 to 4.8 Myr after CAI formation. Some angrites (ADOR, LEW86010) contain no evidence for the former presence of live <sup>26</sup>Al, consistent, with their young Pb-Pb ages. NWA2976 is also characterised by elevated  $\delta^{26}$ Mg\* which, given its anomalous O-Cr isotope ratios compared to all other basaltic meteorites, suggests these small excesses of δ<sup>26</sup>Mg\* in basaltic meteorites reflect decay of <sup>26</sup>Al and are not nucleosynthetic anomalies. (3) Olivines from four main group pallasites yield deficits in  $\delta^{26}$ Mg\* = -0.0157±0.0015%o, which dates olivine crystallization and isolation from high Al/Mg parts of the pallasite parent body at its core-mantle boundary to 0.97±0.10 Myr after CAI formation. Work is ongoing to examine if other meteoritic material (aubrites, ureilites) with sub-chondritic Al/Mg has  $\delta^{26}$ Mg\* deficits that may enable precise dating of differentiation on their parent bodies.

## Predicting solidus temperatures and modes of mantle peridotites

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The peridotite solidus (T<sub>s</sub>) is an important control on the initiation and extent of adiabatic mantle melting. While the pressure (P) - T<sub>s</sub> relationship of fertile peridotite is fairly well constrained [1, 2], the relationship between composition (X) and T<sub>s</sub> is less well quantified. We have developed an algorithm based on mineral-mineral exchange and distribution coefficients (Kds) and stoichiometric and mass balance constraints for calculating the mode and compositions of phases (oliv, opx, cpx, sp, gt) for subsolidus peridotite as a function of T, P, and X. This in turn allows us to estimate the MgO, Na<sub>2</sub>O, and K<sub>2</sub>O (wt %) of near-solidus melt (F = 0.005) using T-, P-, and X-dependent crystal-liquid K<sub>d</sub>s and the assumption that the solidus mode and the mode at F = 0.005are equal. Using 4-phase saturated liquids in CMAS+Fe [3] and CMAS+Na [4], and experimentally determined solidii on natural peridotites [1, 2, references therein, and unpublished data], we parameterized the peridotite solidus as a function of P, and liquid MgO and K<sub>2</sub>O contents. The data set consists of 32 model system experiments and 33 T<sub>s</sub> values from fertile to depleted peridotites; the 65 data points cover a P, T, and Mg# range of 1 to 4.2 GPa, 1210 to 1594°C, and 1 to 0.75, respectively. The mean T deviation of the fit is ~7°C, i.e., less than the uncertainty on any one determination (estimated to be at least  $\pm 10$ -15°C); the largest mismatch is -26°C ( $T_{exp} - T_{calc}$ ).

The calculated T<sub>s</sub> of a primitive mantle (PM) composition [5] varies from 1239°C at 1 GPa to 1506°C at 3.5 GPa, in good agreement with [1] (mean T deviation is ~13°C); over this pressure range the modal opx/cpx ratio at the solidus decreases from 1.24 to 0.04. The calculated solidus for a partially depleted peridotite [6] (MgO = 41.6 wt %), is 45-60°C higher than that calculated for the PM over this P range; the modal opx/cpx ratio at the solidus decreases from 2.88 to 0.97. For a given bulk composition, our calculations show that small changes in silica content have a relatively large effect on opx abundance at the solidus, which in turn affects the P of opx-out along the solidus.

[1] Hirschmann, MM (2000) *G3* 1, #2000GC000070. [2] Herzberg, C *et al.* (2000) *G3* 1, #2000GC000089. [3] Gudfinnsson, GH & Presnall, DC (2000) *J Pet* **41**, 1241-1269. [4] Walter, MJ & Presnall, DC (1994) *J Pet* **35**, 329-359. [5] McDonough, WF & Sun, S-S (1995) *Chem Geol* **120**, 223-253. [6] Wasylenki, LE *et al.* (2003) *J Pet* **44**, 1163-1191.