Mg and Si isotopes in the solar system

R. CHAKRABARTI* AND S.B. JACOBSEN

Dept. of Earth and Planetary Sciences, Harvard University, Cambridge, MA 02138 (rama@eps.harvard.edu)

Chemical heterogeneity in our solar system is welldocumented. Silicate portions of the Earth, Moon and Mars are different from the chondritic meteorites in their volatile element budgets and in Mg/Si. Isotopic compositions also vary between Earth and other Solar System bodies. Chemical and isotopic differences between Earth and bulk meteorites can be broadly attributed to the following processes: (1) Terrestrial fractionation, (2) Fractionation and incomplete mixing in the solar nebula prior to accretion and (3) Late injection of presolar material and its inhomogeneous distribution.

It has been suggested [1] that the Earth and Moon have heavier Si-isotopic compositions compared to meteorites, all of which have indistinguishable Si-isotopes, resulting from core formation on Earth. Several efforts have been initiated to resolve the Mg isotopic compositions of the bulk Earth and meteorites [2, 3] but available results are contradictory.

We have initiated a combined Mg- and Si-isotopic study in terrestrial and planetary samples to address the following questions: (1) What are the BSE Mg- and Si-isotopic compositions? (2) Does the BSE and average chondrites have resolvable differences in Si and Mg-isotopic compositions and (3) if they are different, what is the most plausible explanation? So far our Mg-isotopic data suggest that terrestrial olivines from peridotites and whole-rock basalts show similar Mg-isotopic compositions which are slightly lighter than those of whole-rock chondrites. Olivines from several pallasites show heavier Mg isotopic compositions compared to both chondrites and terrestrial samples. This difference suggests that pallasites, likely derived from the core-mantle boundary of early differentiated planetesimals, sampled Mg-isotopic heterogeneity in the solar nebula.

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In situ SIMS micro-baddeleyite U-Pb dating method for mafic rocks

K.R. CHAMBERLAIN¹*, T.M. HARRISON², A.K. SCHMITT², L.M. HEAMAN³, S.M. SWAPP¹, AND A.K. KHUDOLEY⁴

¹Dept. of Geology and Geophysics, Univ. of Wyoming, Laramie, WY, 82071 USA (*correspondence: kchamber@uwyo.edu)

²Department of Earth and Space Sciences, UCLA, Los Angeles, CA 90095-1567, USA

³Dept. of Earth and Atmospheric Sciences, University of Alberta, Edmonton, AB T6G 2E3, Canada

⁴Geological Department, St. Petersburg State University, University nab. 7/9, St. Petersburg, 199034, Russia

Baddeleyite (ZrO_2) is a favorite target for U-Pb dating of mafic and ultramafic rocks. Although samples with grains that are large enough (100-200 micron long) for traditional separation and dating methods are scarce, micro-baddeleyite crystals (<20 microns long) are nearly ubiquitous. With refined methods, it is possible to separate 10 to 20 micron long grains for ID-TIMS analysis.

We are currently developing an *in situ* technique to date micro-baddeleyite crystals using the CAMECA ims 1270 ion microprobe, avoiding the separation and dissolution steps of ID-TIMS and permitting analysis of even smaller grains (2-5µm). The ion microscope optics of the ims 1270 can reduce the effective sampling diameter of a 20 µm primary ion beam to a few to 10 µm. An aperture blocks undesired secondary ions, in particular common Pb ions emitted from host phases surrounding micro-baddelevite, and the target crystals can be dated in polished thin sections. CAMECA ims 1270 analytical conditions also involve oxygen-flooding the sample chamber ($\sim 10^5$ Torr) to enhance Pb secondary ion yields for baddeleyite by approximately 10-fold, and a UO₂/U vs. U/Pb relative sensitivity calibration. Three baddeleyite standards analyzed under these conditions yielded U/Pb reproducibilities that only marginally exceeded those determined for NIST610 glass, suggesting that crystal orientation related bias in U-Pb sensitivity is minor to absent. Proterozoic-aged standards yielded 2 sigma precisions on ²⁰⁷Pb/²⁰⁶Pb dates from 1% to 0.2%. Preliminary results for micro-baddeleyite from Ordovician and Mesoproterozoic-aged mafic sills analyzed in situ yield U-Pb ages roughly equivalent in precision to those achieved on larger standard baddelevite grains. Precisions on Phanerozoic ²⁰⁶Pb/²³⁸U dates are ~6%; Precambrian ²⁰⁷Pb/²⁰⁶Pb date precisions are 1% or less. The *in* situ technique is applicable to most unmetamorphosed to lowgrade mafic rocks including dikes, sills, fine-grained volcanic rocks and extra-terrestrial rocks, greatly increasing the number of mafic samples that are dateable by U-Pb methods.