

Evidence for an Earth-Moon impact event 800 Ma ago

N.E.B. ZELLNER¹, J.W. DELANO² AND T.D. SWINDLE³

¹Department of Physics, Albion College, Albion, MI 49224, USA (zellner@albion.edu)

²University at Albany (SUNY), Albany, NY 12222, USA (jdelano@atmos.albany.edu)

³University of Arizona, Tucson, AZ 85721, USA (tswindle@U.Arizona.Edu)

Lunar Impact Glasses

Lunar impact glasses offer the potential for providing compositional information about local and remote areas of the Moon [1-3] and may also place constraints on the impact history in the Earth-Moon system.

Previous workers have reported $^{40}\text{Ar}/^{39}\text{Ar}$ ages of ~800 Ma on lunar samples, such as rock fragments (Apollo 12; 4) and impact glasses (Apollo 14, 16; 5, 6). We present additional $^{40}\text{Ar}/^{39}\text{Ar}$ ages on Apollo 14 and 17 impact glasses. When all of these data are viewed collectively, there is a suggestion that there may have been an increase in the impact flux at ~800 Ma.

Results

Of the 62 glasses for which an age has been determined by these authors, 8 have diverse compositions with ages ~800 Ma, as shown in Figure 1. This age is similar to the best current estimate for the age of the Copernicus impact event (~800±15 Ma; [4, 7]).

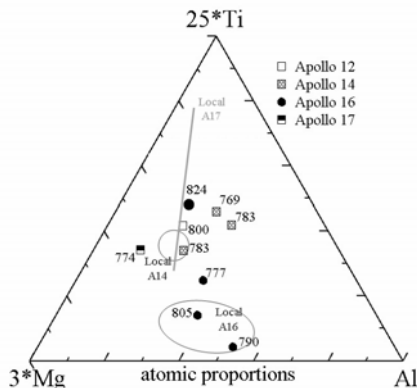


Figure 1; Refractory element compositions of lunar samples (including impact glasses) with ~800 Ma ages.

The higher frequency of impact glasses with this age is consistent with a transient increase in the impact flux at this time. Perhaps the Copernicus impact event was only one of numerous large events (and many small local events) that occurred.

[1] Delano (1991) *GCA* **55**, 3019-3029. [2] Zellner *et al.* (2002) *JGR* **107(E11)** 5102, doi,10.1029/2001JE001800. [3] Delano *et al.* (2007) *MAPS* **42**, 6, 993-1004. [4] Barra *et al.* (2004) *LPSC* **35**, 1365.pdf. [5] Zellner *et al.* (2003) *LPSC* **34**, 1157.pdf. [6] Borchardt *et al.* (1986) *PLPSC* **17th**, p. E43-E54. [7] Bogard *et al.* (1994) *GCA* **58**, 3093-3100.

Early Mesozoic high pressure metamorphism in the Lhasa Block, Tibet: Implications for the growth of the Cimmerian subcontinent

L. ZENG^{1*}, J. LIU-ZENG² AND K. XIE²

¹Institute of Geology, Chinese Academy of Geological Sciences, Beijing 100037, China

(*correspondence: changting1970@yahoo.com)

²Institute of Tibetan Plateau Research, Chinese Academy of Sciences, Beijing 100871, China

The traditional Lhasa Block is part of the Cimmerian subcontinent which was accreted to the southern margin of the Laurentia continent at Mid-Mesozoic [1, 2]. The east-west trending Sumdo-Jiaying eclogite belt, ~200 km east of Lhasa, provides information to constrain how the Cimmerian subcontinent was assembled. This eclogite belt consists of two types of eclogites: kilometers long and several hundreds meters wide massive blocks and sub-meter size subparallel layers or lentoids. Both are enclosed within late Paleozoic quartzite and carbonates. Bulk-rock major and trace element data indicate that these eclogites have preserved an N-MORB type geochemistry [3]. P-T estimations show that they experienced high pressure metamorphism at ~2.6 GPa and ~650°C. Sm-Nd whole rock isochron on 11 data yields an age of 239±16 Ma. Petrography and trace element data suggest that this age represents the timing of high pressure metamorphism. These data show that the Sumdo-Jiaying oceanic block was subducted to a depth of ~75 km and underwent high pressure metamorphism at ~239 Ma. Such an early Mesozoic event of subduction of oceanic crust within the Lhasa Block suggests that (1) at least one oceanic basin existed concurrently with the Paleo-Tethys and (2) the Cimmerian subcontinent was assembled by a number of smaller continental or oceanic blocks that were scattered within the Paleo-Tethyan Ocean.

[1] Dewey, Shackleton, Chang & Yiyin (1988) *Phil. Trans. Roy. Soc. London*, series A, **327**, 379-413. [2] Hsu, Pan, Sengor *et al.* (1995) *Int. Geol. Rev.* **37**, 473-508. [3] Zeng, Liu & Xie (2007) *Eos, Trans. AGU*, **88(47)** Fall Meet. Suppl. Abstract T23D-1637.