

# Geochemistry of geothermal fluids from the Maanping Hotspring Area, Jiangxi Province, SE-China

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## Background

The Maanping Hotspring Area is located in the central part of Jiangxi Province that is one of the Provinces in which hot springs are most widely distributed in China. Strata exposed in the area include the Sinian System (Z) composed of schists, the Cretaceous System (K) characterised by sandstone and conglomerate, and Yanshanian granites ( $\gamma_5^2$ ) as well as Quaternary sediments (Q). Distributed along a NE-trending fault, more than 5 hot springs issue 39°C to 42°C waters in the study area.

## Hydrochemistry and subsurface temperature estimation

The hot springs are Na-K-HCO<sub>3</sub> waters with TDS higher than 1000 mg/L. The geothermal reservoir temperatures are estimated to be from 72°C to 82°C by silica geothermometers. The fluid-mineral equilibria studies show that chalcedony, calcite, aragonite and dolomite are close to equilibrium with solution in the geothermal system.

## Isotope geochemistry

The chemical composition of the geothermal gases collected from the bubbling hot springs in the area using the water-replacement method are the carbon dioxide-dominant type with CO<sub>2</sub> concentrations ranging from 97.76% to 97.90%, N<sub>2</sub> from 1.75% to 1.88%, Ar from 0.025% to 0.027%, and O<sub>2</sub> from 0.32% to 0.33% respectively.

The  $\delta^{13}\text{C}$  values of CO<sub>2</sub> for the geothermal gases are from -4.18‰ to -7.0‰ (PDB) showing that the CO<sub>2</sub> may come from the mantle source, thermogenic decomposition of carbonates in deep crust, or the mixture of the two components.

The  $^3\text{He}/^4\text{He}$  ratios of the geothermal gases range from 0.15 Ra to 0.18 Ra suggesting input of radiogenic  $^4\text{He}$  in crystalline rocks of the crust in the area.

The  $\delta^{18}\text{O}$  and  $\delta\text{D}$  values of the geothermal waters in the hot spring area vary from -6.84‰ to -7.23‰ (SMOW) and from -44.1‰ to -48.8‰ (SMOW) respectively. In comparison, the values of local shallow groundwaters range from -3.97‰ to -5.46‰ for  $\delta^{18}\text{O}$  and from -22.3‰ to -37.6‰ for  $\delta\text{D}$ . The evidence shows that the geothermal waters are recharged from a higher altitude area compared to local shallow groundwaters.

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