Martian subsurface biomes: how detectable are they?

T.C. ONSTOTT¹, L. M. PRATT², S. M. CLIFFORD³, B. SHERWOOD LOLLAR⁴ And T. J. PHELPS⁵

¹ Princeton University, Princeton, NJ, USA; tullis@princeton.edu

² Indiana University, Bloomington, IN, USA; prattl@indiana.edu

³Lunar Planetary Institute, Houston, TX, USA; Clifford@lpi.usra.edu

⁴University of Toronto, Toronto, ON, CANADA;

bslollar@chem.utoronto.ca

⁵Oak Ridge National Laboratory, Knoxville, TN, USA;

phelpstj@ornl.gov

Prior papers examining the habitability of Mars or life on Mars originally focused on surface dwelling photosynthetic or chemolithotrophic microorganisms but over the last ten years have emphasized subpermafrost ecologies. Many papers derive biomass concentrations based upon a bulk energy fluence estimates (Jg⁻³s⁻¹) divided by an assumed maintenance energy flux (Jcell⁻¹s⁻¹). Derived on the assumption that oxidants are entirely derived from the atmosphere or oxidized surface rocks and in a few papers that reductants, H₂ and CO, are derived from the atmosphere and that they are transported to the subsurface by gas diffusion, basal melting of the cryosphere and hydrothermal fluid circulation. Radiolysis of ice and water, however, provides a subsurface source of H₂, H₂O₂, OH and O₂ at a rate of nM yr ¹, which is greater than the surface fluence and keeps the subsurface biosphere from approaching ergonic death [1]. The potential energy flux can be 10⁴⁻⁵ greater if it is concentrated in highly fractured and porous rock, e.g. beneath impact structures. This rate of 10^{-11} Jg⁻³s⁻¹ is comparable to that estimated for hydrothermal activity and weathering reactions. Anaerobic metabolites, e.g. CH₄, N₂ and H₂S, will be trapped as hydrates in a thickening cryosphere; whereas He and H₂, will diffuse through the permafrost/cryopeg zone. Within the subphotic dissociation zone the upper surface of the cryosphere evaporates, the hydrates exsolve and the trapped gases will also diffuse to the surface. At the interface where H₂, CH₄ and H₂S encounters O₂ diffusing downward from the Martian atmosphere, oxidation of reduced gas phases could produce acidic solutions that could amend subsurface brine volumes by melting the permafrost. Terrestrial analogs of this interface should exist and merit investigation as they are more accessible to near future robotic missions than the deep subsurface of Mars.

References

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