The existence and mobility of hydronium ion in the structure of jarosite

JURAJ MAJZLAN¹, ULLA GRO NIELSEN² AND CLARE P. GREY²

¹Institute of Mineralogy, Petrology and Geochemistry, Albert-Ludwig-University of Freiburg, Albertstrasse 23b, Freiburg, D-79104 Germany (Juraj.Majzlan@minpet.uni-freiburg.de)

²Department of Chemistry and Center for Environmental Molecular Science, State University of New York, Stony Brook NY 11794-3400, USA

Jarosite is a common mineral of acid mine drainage systems. The most common end-members are jarosite *s.s.* $(KFe_3(SO_4)_2(OH)_6)$, natrojarosite $(NaFe_3(SO_4)_2(OH)_6)$, and hydronium jarosite $((H_3O)Fe_3(SO_4)_2(OH)_6)$. Despite its ubiquitous presence as a component of the jarosite solid solution, the properties of hydronium jarosite are not known well. In particular, the question of the existence of the hydronium ion in this structure has not been resolved yet.

The structure of jarosite is built by triangular sheets of octahedra $Fe(OH)_6$, decorated by sulfate tetrahedra. The Fe^{3+} sites are partially (5-15 %) vacant. The monovalent cations reside in cavities between the sheets.

We have investigated a series of jarosite samples and their deuterated analogs by ²H magic-angle spinning nuclear magnetic resonance (MAS NMR) and quasi-elastic neutron scattering (QENS) on ¹H atoms.

²H MAS NMR results clearly distinguish hydroxyls (OD) in the structure from scarce OD₂ groups bound to iron and abundant OD_n (n = 2 or 3) groups not bound to iron. The δ_{iso} values for these three sites are $\approx 0, \approx 70$, and ≈ 260 ppm, respectively. The OD₂ groups bound to iron correspond to water molecules that coordinate vacant sites in the structure. The OD_n groups are either OD₂ or OD₃ molecules in the cavities. Integration of the intensities of the peaks suggests that n = 3, i.e., the cavity is occupied by a D3O+ group. The shape of the NMR spectra indicates that the D3O+ deuterons are highly mobile and tumble in their cavities.

QENS spectra show that the hydrogen in the sample is in motion already at \sim 17 K. The quasi-elastic line is broadened when compared to the vanadium standard. The spectra collected at 50, 120, 170, 220, and 300 K show further broadening of the quasi-elastic line. Separation of the signal due to rotation and possible translation of the hydrogen atoms in the QENS spectra is now in progress.

Jarosites in acid mine drainage environments: Formation, mineralogy, stability

K.A. HUDSON-EDWARDS¹, A.M.L. SMITH^{1,2,3}, W.E. DUBBIN² AND K. WRIGHT³

¹Research School of Earth Sciences at UCL-Birkbeck, Gower Street, London, WC1E 6BT, UK (k.hudson-edwards@bbk.ac.uk)

²Department of Mineralogy, The Natural History Museum, Cromwell Road, London SW7 5BD, UK

³Davy Faraday Research Laboratory, The Royal Institution of Great Britain, 21 Albemarle St., London W1S 4BS, UK

Jarosites form a subgroup of the alunite supergroup of minerals, and have a general formula of $AB_3(TO_4)_2(OH)_6$, where A represents cations with a coordination number greater than or equal to 9 (K⁺, Na⁺, NH₄⁺, H₃O⁺, Ag⁺, Tl⁺, Pb²⁺), B represents cation sites with octahedral coordination (mainly Fe³⁺, but also Cu²⁺), and T represents cations with tetrahedral coordination (SO₄, AsO₄) (Jambor 1999). Jarosites are commonly found in ferric-rich, acidic (pH < 3), oxic, acid mine drainage (AMD) environments, where they are important stores of SO₄ and of potentially toxic elements such as Pb and As.

In AMD environments such as the Río Tinto, Spain, and the Mathiatis mine. Cyprus, plumbojarosite [Pb_{0.5}Fe₃(SO₄)₂(OH)₆] and jarosite itself [KFe₃(SO₄)₂(OH)₆] are common, and form as a result of the decomposition of pyrite or precipitation of sulphate-rich, ferric solutions. Batch experiments and computational modelling, carried out to investigate the stability of jarosite, plumbojarosite and beudantite [PbFe₃(SO₄)(AsO₄)(OH)₆] under acidic (pH 2) and alkaline (pH 8) conditions, have shown that these minerals all dissolve incongruently. The pH 2 dissolution of jarosite yields excess K and SO₄ relative to Fe in solution, which may be related to selective dissolution of the A- (K) and T- (SO₄) sites, since Fe is located deep within the T-O-T jarosite structure. Of the pH 2 dissolutions, only that of beudantite yields a secondary, Fe hydroxide-like phase.

The alkaline (pH 8) dissolutions all yield secondary phases: jarosite yields goethite, plumbjarosite and beudantite yield Fe hydroxide, and beudantite also yields $PbSO_4$. There is evidence for some co-precipitation and resorption of the Pb and As in and onto the newly-formed phases, suggesting that the remainder of these potentially toxic elements may be released back to AMD waters.

Reference

Jambor, J. (1999) Can. Miner. 37, 1323-1341.