

**1.1.P02****Water partition coefficients between nominally anhydrous minerals and basaltic melts**C. AUBAUD<sup>1</sup> E.H. HAURI<sup>2</sup> AND M.M. HIRSCHMANN<sup>1</sup><sup>1</sup>Dept. Geology & Geophysics, Univ. Minnesota, Minneapolis, MN55455 (aubau001@umn.edu)<sup>2</sup>DTM, Carnegie Institute of Washington, Washington DC

Partitioning of water between peridotite minerals and basaltic magma has a significant influence on the initiation of melting in the mantle and on the rheological structure of the lithosphere. To investigate mineral/melt and mineral/mineral partitioning of H<sub>2</sub>O applicable to the mantle, we have conducted multiple saturation experiments consisting of hydrous basalt+ol±opx±cpx in a piston-cylinder apparatus at pressures of 1-2 GPa, temperatures of 1230-1300°C and bulk initial water contents of 4 to 6 wt.%. Minerals are similar in composition to those expected in spinel peridotite. We measured H<sub>2</sub>O in melts and minerals (ol, opx, cpx) by SIMS using methods described by [1].

Resulting liquids have 3.6-10 wt.% H<sub>2</sub>O and average mineral/melt partition coefficients as follows:  $D^{ol-melt}=0.0012\pm0.0003$  (n=10),  $D^{opx-melt}=0.014\pm0.006$  (n=9), and  $D^{cpx-melt}=0.018\pm0.008$  (n=3). Mineral/mineral partition coefficients are  $D^{ol-opx}=0.10\pm0.03$  (n=6),  $D^{ol-cpx}=0.08\pm0.02$  (n=3) and  $D^{cpx-opx}=1.3\pm0.1$  (n=2). Observed partition coefficients are reproducible between experiments and systematic variations with pressure, temperature or concentration of H<sub>2</sub>O are not apparent, possibly because of the relatively small range of pressures and compositions examined. For a peridotite consisting of 58% ol, 30% opx, 10% cpx, and 2% spinel (assumed nominally anhydrous) the calculated bulk sol-liq  $D$  is  $0.007\pm0.002$  confirming that water is highly incompatible in mantle minerals.

Compared to conventional trace elements, water has a behavior between that of La and Ce, in accordance with studies on natural basaltic glasses (e.g. [2]). If this bulk  $D$  is applicable to the deeper parts of the MORB melting regime, then following [3], we can estimate the effect of H<sub>2</sub>O on peridotite partial melting: for a water concentration of the source of 200 ppm, the near-solidus melt would contain 2.9 wt.% water. For a  $\Delta H_{fusion}$  of 50 kJ/mol, the freezing point depression is ~40°C, which corresponds to initiation of melting 15 km deeper than the anhydrous solidus, somewhat less than previous estimates (~50 km, [4]).

**References**

- [1] Koga et al. (2003) *G3* **4**, 1-20.
- [2] Dixon et al. (1988) *EPSL* **90**, 87-104.
- [3] Hirschmann et al. (1999) *J. Petrol.* **40**, 831-851.
- [4] Hirth and Kohlstedt (1996) *EPSL* **144**, 93-108.

**1.1.P04****Solids of alkaline aluminosilicate system under H<sub>2</sub>O/H<sub>2</sub>-fluid pressure (P<sub>total</sub>=2 kbar)**

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The vapor-saturated solids of NaAlSi<sub>3</sub>O<sub>8</sub>-H<sub>2</sub>O-H<sub>2</sub>, NaAlSi<sub>3</sub>O<sub>8</sub>-SiO<sub>2</sub>-H<sub>2</sub>O-H<sub>2</sub>, and NaAlSi<sub>3</sub>O<sub>8</sub>-KAlSi<sub>3</sub>O<sub>8</sub>-SiO<sub>2</sub>-H<sub>2</sub>O-H<sub>2</sub> systems at 2 kbar over the entire range of gas phase composition from pure water to pure hydrogen were studied in the internally heated gas-media pressure vessel. In the experiments, various H<sub>2</sub>O/H<sub>2</sub> compositions were controlled directly, rather than using a solid media buffers.

At addition of hydrogen to the considered H<sub>2</sub>O-saturated systems a pronounced temperature minimum on solidus curves appears at  $f(H_2)=300-500$  bars. The solidus temperatures decrease concerning H<sub>2</sub>O-saturated systems on 22K for Ab, 30K for Ab-Qz, and 40K for haplogranite (Ab-Or-Qz) system. At the further increase of  $f(H_2)$  in H<sub>2</sub>O-H<sub>2</sub> fluid the solidus temperatures increase reaching the maximum at pressure of pure hydrogen.

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For an estimation of hydrogen influence on formation granite melts and comparison with published data received with use of buffer reactions in system Fe-O the MW and WI buffers which use as indicators of  $f(O_2)$  in natural and experimental processes have been investigated by the same technique. Experimental researches show that in the H<sub>2</sub>O presence at MW and WI buffers the high values of  $f(H_2)$  are reached. Hydrogen reacts with wustite and iron and the mixed compounds form. The monovariant buffer reactions are transformed in the divariant fields. In the result, the stability of widely widespread in granites magnetite is considerably shifted to the reducing area.